

Graptolite biostratigraphy of the Shihtien Formation (Darriwilian) in West Yunnan, China

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A rich mid-Ordovician graptolite fauna is reported from the Shihtien Formation at Baoshan and Shidian in West Yunnan Province, SW China. The fauna comprises *Didymograptus artus*, *D. murchisoni*, *D. spinulosus*, *Pterograptus* sp., *Hustedograptus vikarbyensis*, *H. teretusculus*, *Archiclimacograptus angulatus*, *A. riddellensis*, *Haddingograptus oliveri*, *Proclimacograptus angustatus*, and dendroids such as *Dictyonema*, *Ptilograptus*, *Dendrograptus*, *Callograptus*, indicating an age of mid- to late Darriwilian. Based on the fauna, two graptolite assemblages are recognized, in ascending order, the *Didymograptus artus* and *Didymograptus murchisoni* biozones. The graptolite fauna shows a considerable similarity to those contemporary in Baltica and the Yangtze Region of South China. • Key words: graptolites, biostratigraphy, Darriwilian, Ordovician, West Yunnan, China.

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West Yunnan has been widely regarded as a constituent of the Sibumasu Terrane (sometimes called the Shan-Thai Terrane or Paleopllate) during the Ordovician (Chen & Rong 1992; Fortey & Cocks 1998, 2003; Metcalfe 1998). The terrane embraces West Yunnan, East Burma, most of the Malay Peninsula and Sumatra, and is bounded to the east in Yunnan by the Changning-Menglian Suture (Fang *et al.* 1992, Liu *et al.* 1993, Metcalfe 1998). The affinity of the Sibumasu Terrane with other blocks or continents, such as South China, North China, and Gondwana, remains uncertain, partly because of poor fossil discoveries in Sibumasu (see Zhou & Zhen 2008). Fortey & Cocks (1998, 2003) concluded that the carbonate succession of the Cambrian-Ordovician transition interval in the terrane displays strong similarity to that in North China and central Australia, whereas those trilobite faunas of the later Ordovician are nearly identical with those of South China, despite the fact that Caradoc brachiopods from Burma show a modest level of endemism (Cocks & Zhan 1998).

The Ordovician rocks of West Yunnan have been poorly known until now. Though reported present in many localities, the Ordovician rocks were strongly affected by tectonic activities of many phases, especially the recent Himalayan that commenced in the Palaeogene. The earliest geological work in the region were carried out by L.V.

Loczy in 1880, and J.C. Brown in 1908–1910. Loczy (1893) reported the first discoveries of Ordovician trilobites from the region, and Brown (1913, 1916) published a primary introduction to the Ordovician and Silurian rocks. Abundant fossils collected during Brown's investigation, including cystoids, crinoids, brachiopods, gastropods, cephalopods, trilobites and graptolites, were subsequently published by Reed (1917); the graptolites, mostly from Silurian rocks, were identified by Gertrude Elles (*in* Reed 1917) as of Llandovery age. Yin & Lu (1937) collected and described the Jenhochiao section (now called Banpo) in Shihtien (modern pinyin transliteration, Shidian). They reported the discovery of Ordovician brachiopods, trilobites, cephalopods and graptolites including *Didymograptus nanus*, *Didymograptus* sp. indet., *Dictyonema* sp., and *Desmograptus?* sp., which were described briefly and illustrated by Yin (1937). A subsequent investigation of the Baoshan region in West Yunnan by Sun in 1938 resulted in a tripartite subdivision of the Ordovician rocks (Sun 1945): the Shihtien Formation (Arenigian and Llanvirnian, in which *Didymograptus nanus* and *Didymograptus murchisoni* were reported); the Lower Hungshuitang Formation (Llandeilian); and the Upper Hungshuitang Formation (Caradocian). Yu (1962) described some additional graptolite species that he assigned to *Ptilograptus* and a new genus *Zigzagigraptus*.

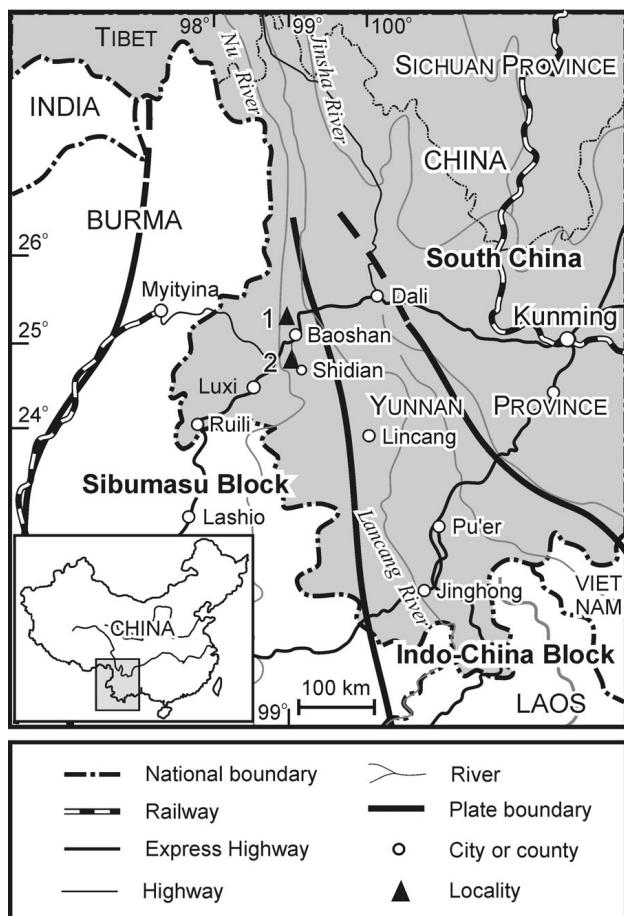


Figure 1. Location of the Laojianshan section in Baoshan, and the Banpo section in Shidian, West Yunnan, China.

Lithological sequence of the Ordovician strata

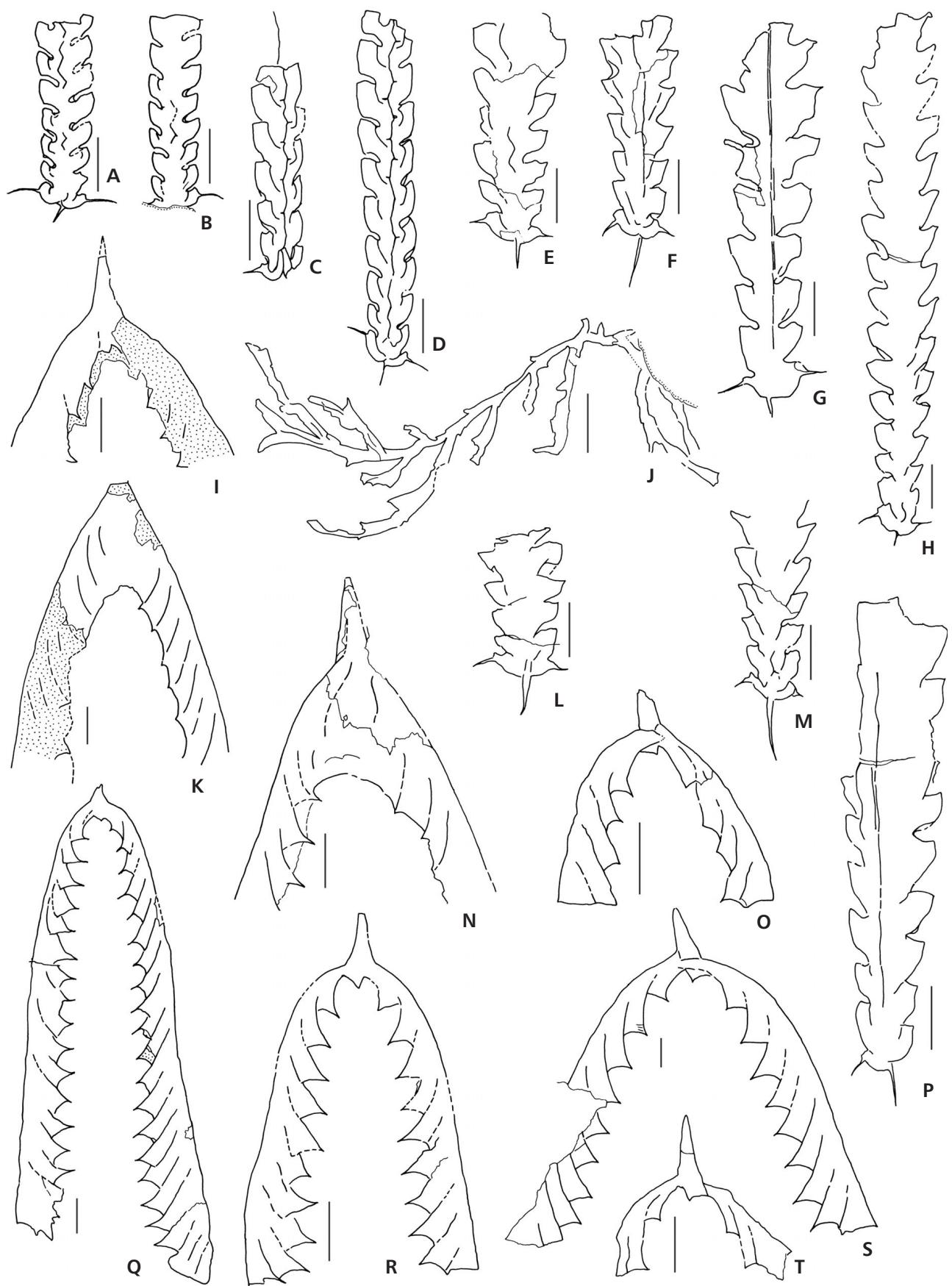
We have recovered a more diverse Ordovician graptolite fauna from both the Banpo section in Shidian and the Laojianshan section in Baoshan than has previously been reported from the region (Fig. 1). The Laojianshan section ($N\ 25^{\circ}09.126'$, $E\ 99^{\circ}02.975'$), which we sampled in 2003 and 2007, is located at Laojianshan Mountain near Upper Yanqing village, ca 17 km to the northwest of downtown

Baoshan City. The section, exposed along the country road from Zhugeyin Town to Yangliu Town, includes a complete sequence of Ordovician strata that, in ascending order, comprise the Yanqing, Mantang, Laojianshan, Shih-tien, and Pupiao formations (see Yunnan Geological Survey 1980).

The Yanqing Formation is composed of sandstone with a few layers of laminated limestone and siltstone, and contains trilobites, cystoids and conodonts, indicating an early Tremadocian age. The Mantang Formation consists of sandstone and siltstone, and is regarded as late Tremadocian in age, based on a few brachiopods and graptolites of *Rhabdinopora flabelliformis*, which have been reported from other nearby sections (Yunnan Geological Survey 1980). The Laojianshan Formation is 743 m thick, and typified by purple-red and greenish-grey sandstone, siltstone and shale, with developed fine lamination, cross-bedded and graded sequences. The interbedded greenish-grey shale contains a few trilobites and graptolites (*Callograptus* and *Didymograptus*). More graptolites were found in other sections nearby, and based on these, three graptolite biozones have been recognized, in ascending order *Acanthograptus*, *Didymograptus protobifidus*, and *Undulograptus austrodentatus* biozones (Yunnan Geological Survey 1980), suggesting an age spanning Floian to earliest Darriwilian for the formation. However, the graptolites are not yet described.

The Shihtien Formation overlies conformably the Laojianshan Formation and underlies conformably the Pupiao Formation. It is typified by greenish-grey and green-yellow siltstone and shale, with a few layers of thin-bedded micrites, containing fairly abundant graptolites, trilobites and brachiopods of Darriwilian age. The formation was derived from the original 'Shihtien Beds' of Brown (1916), which was later split by Yin & Lu (1937) into a 'Shihtien bed' (lower) and a 'Hengshuitang Limestone' (upper). The 'Shihtien bed' was further redefined as the 'Shihtien Formation', which refers to the lithology and interval as we use them here, whereas the 'Hengshuitang Limestone' was renamed as the Pupiao Formation (Yunnan Geological Survey 1980). The Pupiao Formation is typified by purple-red and greenish-grey siltstone, shale and micrite, with

Figure 2. Illustration of the Darriwilian graptolites from the Shihtien Formation, mid-Ordovician, West Yunnan. • A, B – *Archiclimacograptus riddellensis* (Harris, 1924); A – NIGP149274 (BSSD-4-8), B – NIGP 149275 (BSSD-4-24). • C, D – *Archiclimacograptus angulatus* (Bulman, 1953); C – NIGP 149279 (BSSD-4-21a), D – NIGP 149277 (BSSD-4-21b). • E, F, L – *Hustedograptus vikarbyensis* (Jaanusson, 1960); E – NIGP 149261 (BPSD-1-19), F – NIGP 149259 (BSSD-4-38), L – NIGP149260 (BSSD-4-58). • G, H, M – *Hustedograptus teretiusculus* (Hisinger, 1840); G – NIGP149272 (BPSD-1-46), H – NIGP149270 (BPSD-1-55), M – NIGP149273 (BSSD-4-37). • I, K, N, Q – *Didymograptus murchisoni* (Beck, 1839); I – proximal part of mature specimen, NIGP149281 (BPSD-1-32); K – proximal part of mature specimen with developed membrane, NIGP149282 (BPSD-1-15); N – proximal part of mature specimen with typical *pakrianus*-like membrane, NIGP149256 (BPSD-1-18); Q – NIGP149255 (BSSD-4-40). • J – *Pterograptus* sp.; NIGP149239 (BPSD-1-25). • O, R-T – *Didymograptus artus* Elles & Wood, 1901; O – proximal part showing the origination of th₂¹ from th₁¹, NIGP149283 (BSSD-3-24); R – NIGP149246 (BSSD-4-8), S – NIGP149284 (BSSD-4-1); T – juvenile specimen showing slender proximal part, NIGP149240 (BPSD-1-3). • P – *Eoglyptograptus dentatus* (Brongniart) *sensu* Bulman, 1963, NIGP149264 (BSSD-4-21). All the scale bars are 1 mm in length. BSSD-samples at the Laojianshan section, BPSD-samples at the Banpo section.



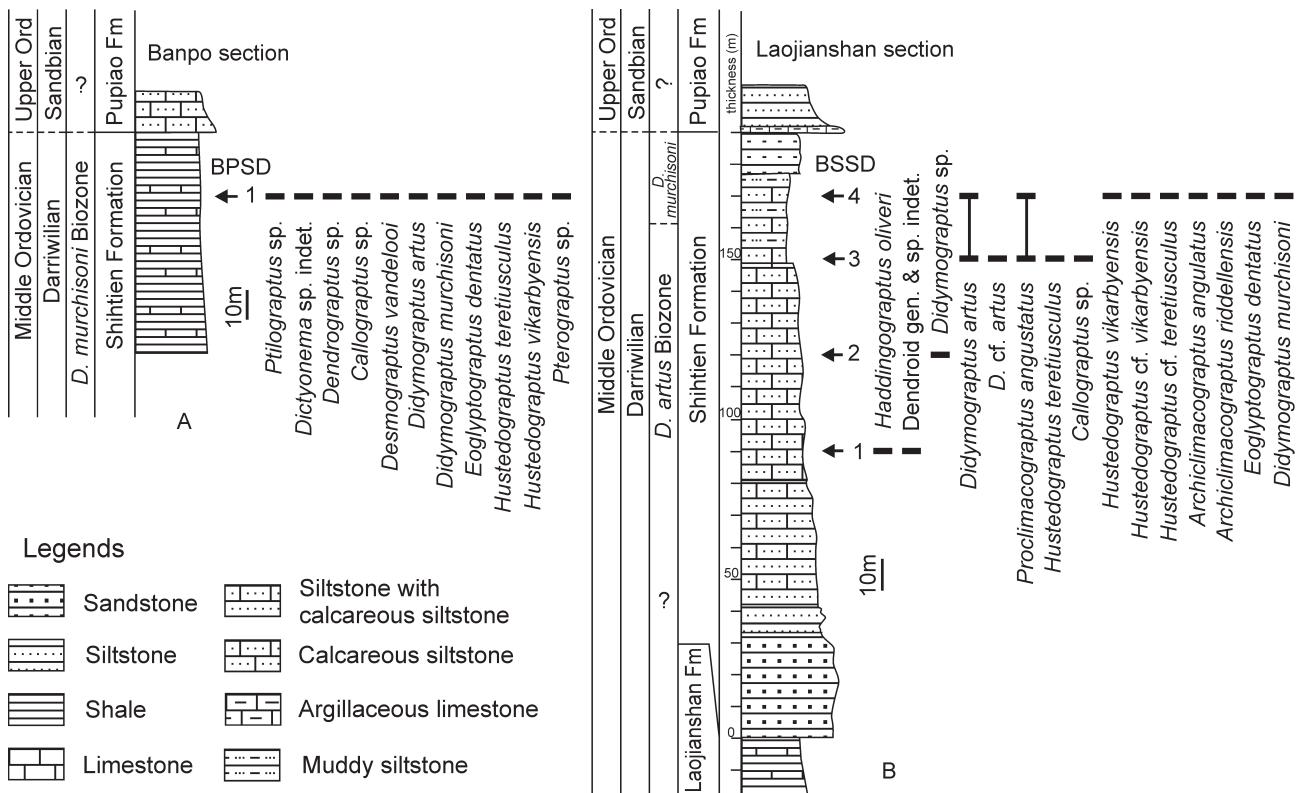


Figure 3. Range chart of the Darriwilian graptolites in the Laojianshan and Banpo sections, West Yunnan.

its base defined at the first appearance of purple-red layers, and its top at the appearance of black shale and chert (Jennochiao Formation, Llandovery).

The Banpo section ($N\ 24^{\circ}46'$, $E\ 99^{\circ}09'$), which we sampled in 2007, is located at Banpo village of Renheqiao (formerly spelt as Jennochiao) Town, ca 6 km to the north of Shidian County town (and ca 60 km from Baoshan). The section includes the Shihtien, Pupiao and Jennochiao formations, and is moderately affected by tectonics. In the upper part of the Shihtien Formation, we discovered the rich Darriwilian graptolite fauna illustrated here (Fig. 3).

Graptolite biostratigraphy and correlations

Elles identified the graptolites collected by Brown in his pioneer expedition as including *Didymograptus murchisoni*, *D. murchisoni geminus*, *D. indentus* and *Climacograptus cf. scharenbergi*, which suggested an age of the *Didymograptus murchisoni* Biozone (Brown 1916, Reed 1917). The rocks bearing these graptolites were assigned by Brown into the basal part of the ‘Pupiao Beds’, but later included into the upper part of the redefined ‘Shihtien Formation’ by Yin & Lu (1937). The latter found additional *Didymograptus nanus* and some dendroids from the rocks, which indicated an earliest Llanvirn age of the ‘*Didymograptus bifidus*’ Biozone (= *D. artus* Biozone).

Our collection includes a more diverse fauna of graptolites (Figs 2, 3), including *Didymograptus artus*, *D. murchisoni*, *D. spinulosus*, *Pterograptus* sp., *Hustedograptus vikarbyensis*, *H. teretiusculus*, *Archiclimacograptus angulatus*, *A. riddellensis*, *Haddingograptus oliveri*, *Proclimacograptus angustatus*, and dendroids such as *Dictyonema*, *Ptilograptus*, *Dendrograptus*, *Callograptus*. The fauna is fairly similar to those found in Norway (Berry 1964, Maletz 1997) and the Yangtze Region of South China (Zhang *et al.* 2007). Based on the occurrences of the fauna, two graptolite biozones, in ascending order the *Didymograptus artus* Biozone and the *Didymograptus murchisoni* Biozone, are recognized herein (Fig. 3).

Didymograptus artus Biozone

The abundance biozone is recognized by the predominant occurrences of the eponymous species, together with *Haddingograptus oliveri*, *Proclimacograptus angustatus*, *Hustedograptus teretiusculus*, uncertain species of *Didymograptus*, and some dendroids (Fig. 2). These graptolites occur mainly in the upper part of the Shihtien Formation. Sample BSSD-3 yielded abundant *D. artus* and associated graptolites, as well as rich trilobites (e.g., *Birmanites yunnanensis* and *Illaenus* sp.). Samples BSSD-1 and BSSD-2

contained relatively fewer specimens of graptolites and trilobites, of which a few specimens are identified as *Didymograptus* sp. but no unequivocal *D. artus* can be confirmed. Therefore, the base of the *D. artus* Biozone is uncertain. The lower part of the Shihtien Formation, composed of sandstone and calcareous sandstone, contains no graptolites. However, *Undulograptus austrodentatus* was found from the top part of the underlying Laojiangshan Formation in some other sections near Shidian (Yunnan Geological Survey 1980, p. 33), and hence the lower part of the Shihtien Formation probably corresponds to the transition interval between the *U. austrodentatus* and *D. artus* biozones, equivalent to the *Undulograptus intersitus* Biozone of the Yangtze Region (South China, Zhang & Chen 2003, Zhang *et al.* 2007) and Australasia (VandenBerg & Cooper 1992), and the *Undulograptus dentatus* Biozone of North America (Maletz 2005). The *D. artus* Biozone in west Yunnan is comparable with that recognized in Chongning, Sichuan, South China (Zhang *et al.* 2007), and in Britain (Fortey *et al.* 1995, 2000). It may also correspond to the *Nicholsonograptus fasciculatus* Biozone and the lower part of *Pterograptus elegans* Biozone in the Jiangnan region of South China (Zhang *et al.* 2007) and in Baltica (Maletz 1997).

Didymograptus murchisoni Biozone

This abundance biozone is characterized by a proliferation of the eponymous species, together with common biserials, including *Hustedograptus teretiusculus*, *H. vikarbyensis*, *Archiclimacograptus riddellensis*, *Archicl. angulatus*, *Eoglyptograptus dentatus*, rare *Pterograptus* sp. and some dendroids. *Didymograptus artus* and *Proclimacograptus angustatus*, which first appear in the underlying *D. artus* Biozone, also extend into this biozone. *Hustedograptus teretiusculus* is fairly common in this biozone, but the species has been proved as long-ranging in both Baltica and South China, and thence is biostratigraphically less significant. The species *Hustedograptus vikarbyensis*, *Archiclimacograptus riddellensis* and *A. angulatus*, as well as *Eoglyptograptus dentatus* (*sensu* Bulman), all first appear in the upper *Pterograptus elegans* Biozone of Norway (Maletz 1997), indicating a correlation of the *D. murchisoni* Biozone in West Yunnan to the upper *P. elegans* Biozone in Baltica. The occurrence of *Pterograptus* sp., although with only one specimen in BPSD-1, further supports this correlation. The *D. murchisoni* Biozone is also well correlated with the same biozone in Britain (Fortey *et al.* 1995, 2000), the Yangtze region of South China (Zhang *et al.* 2007), the lower part of the *Archiclimacograptus riddellensis* Biozone in Australia (VandenBerg & Cooper 1992), and the *Pterograptus elegans* Biozone in Argentina (Brussa *et al.* 2003a, b).

No graptolites have been found from the top part of the Shihtien Formation and the entire Pupiao Formation in the region, and thus no graptolite zones younger than the *Didymograptus murchisoni* Biozone can be identified.

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