

The illaenid trilobites *Vysocania* (Vaněk & Vokáč, 1997) and *Octillaenus* (Barrande, 1846) from the Upper Ordovician of the Czech Republic, Portugal, Spain and Morocco

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The illaenid trilobite *Vysocania* is widely represented in the Upper Ordovician of the Czech Republic, Portugal and Spain, and is one of the most characteristic taxa in the high-latitude peri-Gondwana palaeobiogeographical region to which it is endemic. Three species of the genus that were formerly assigned to *Stenopareia* and *Zetillaenus*, from the Czech Republic and Spain respectively, are herein revised, and a new species is described from the Czech Republic. The genus *Octillaenus*, of which a new species is described from the Upper Ordovician of Morocco, is considered to have arisen from *Vysocania* through pedomorphic (neotenic) processes. The relationships between *Vysocania*, *Octillaenus*, *Stenopareia* and *Zetillaenus* are reviewed and the criteria used in the definition of these taxa discussed. Species assigned to *Vysocania* are *V. vaneki* (type-species), *V. iberica*, *V. moravecii* sp. nov. and *V. panderi*; species assigned to *Octillaenus* are *O. hisingeri* (type-species) and *O. marocanus* sp. nov. *Vysocania* appears for the first time in the lower Berounian (Sandbian) and has its last undoubted record in the upper Berounian (upper Katian); *Octillaenus* is presently restricted to the Kralodvorian (upper Katian). • Key words: *Stenopareia*, *Zetillaenus*, Illaenidae, high-latitude peri-Gondwana, Dalmanitoidean Realm.

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The smooth exoskeleton of illaenids causes considerable difficulties in taxonomic differentiation within this group (e.g. Jaanusson 1954, Lane & Thomas 1983). Numerous Upper Ordovician illaenid species have traditionally been included in the widely conceived genus *Stenopareia* Holm, 1886, which has long served as a ‘waste-bin taxon’ to group illaenid morphologies characterized especially by nine thoracic segments, a highly convex cephalon, glabella defined by short (exsag.) axial furrows and pygidium semicircular in outline. As presently conceived, *Stenopareia* is a polyphy-

letic taxon in need of revision using modern phylogenetic analysis. Such an analysis has proved its usefulness with closely related taxa; e.g. *Bumastoides* Whittington, 1954, analyzed by Carlucci *et al.* (2012). The need for such a revision is common to many Upper Ordovician illaenids.

During the last 70 years, reports of *Stenopareia* species from the high latitude peri-Gondwanan region remained taxonomically stable (e.g. Jaanusson 1954, Šnajdr 1957, Bruthansová 2003). However, crucial morphological differences distinguish these species from Baltic *Stenopareia s.s.*

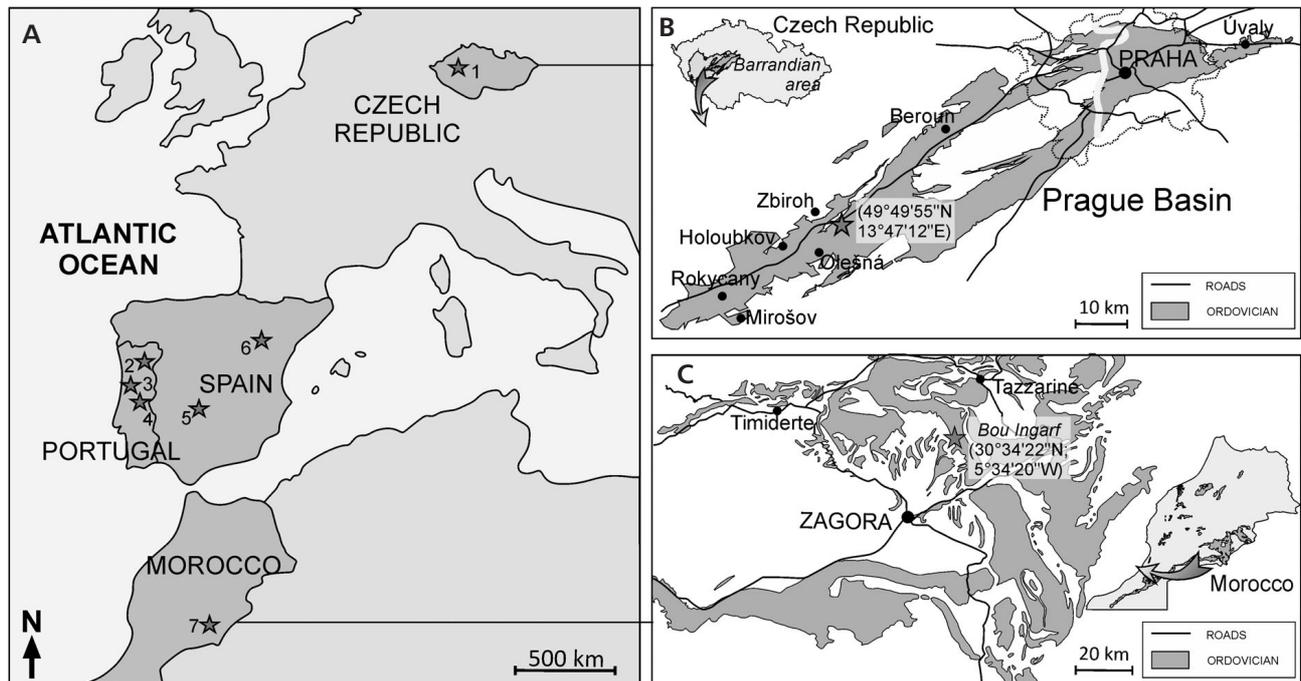


Figure 1. Geographical location of the localities (A) and type-localities for the new taxa (B, C) of *Vysocania* Vaněk & Vokáč, 1997. • A – Czech Republic: 1 – Prague Basin; Portugal: 2 – Torre de Moncorvo, 3 – Buçaco-Penacova, 4 – Mação; Spain: 5 – Ciudad Real, 6 – Zaragoza; Morocco: 7 – Zagora. • B – detail of the Zbiroh type locality for *Vysocania moravecii* sp. nov. • C – detail of the Bou Ingarf type locality for *Octillaenus maroccanus* sp. nov.

These differences, detailed in this work, justify the separation of the peri-Gondwanan species in a different genus. Vaněk & Vokáč (1997) erected the genus *Vysocania* to include only *Zbirovia vaneki* Šnajdr, 1958, which is congeneric with the Czech species previously assigned to the genus *Stenopareia*. For this reason, the genus *Vysocania* should be used to embrace those species. On the other hand, the genus *Octillaenus* Salter, 1867 remained monospecific since its definition, the type species *Octillaenus hisingeri* (Barrande, 1846) being restricted to the upper Katian of the Czech Republic. A new species from a coeval unit of Morocco is described and it is suggested that *Octillaenus* developed from *Vysocania* through pedomorphic processes.

Vysocania is widely represented in the peri-Gondwanan region, being one of the most characteristic trilobite genera in the Upper Ordovician sequences of the Czech Republic, Portugal and Spain (Fig. 1A). *Vysocania* is particularly abundant at the base of the successions bounded by ironstone horizons, in which trilobite diversity is fairly low. *Octillaenus* is currently known from the Czech Republic and Morocco.

Geographical and geological settings of material

Czech Republic. – *Vysocania* is known from several stratigraphical units (Fig. 2) and localities in the Prague Basin of

the Barrandian area (locality 1, Fig. 1A; see Havlíček 1998, Fatka *et al.* 2013 and Kraft *et al.* 2015 for information on the Ordovician stratigraphy of the basin). The specimens of *Vysocania moravecii* sp. nov. come from the lower Berounian Libeň and Letná formations. The specimens of *V. panderi* (Barrande, 1852) originate from the middle Berounian Vinice and Zahofany formations. The base of the Vinice Formation is marked by a discontinuous ironstone ore layer named the Zdice-Nučice ‘Horizon’, characterized by a low diversity trilobite assemblage in which *Vysocania* is one of the most abundant elements. The specimens of *V. vaneki* (Šnajdr, 1958) come from the upper Berounian Bohdalec Formation, the type horizon of the species being the discontinuous, sideritic Karlík ‘Ore Horizon’ at the base of the formation. The specimens of *Octillaenus hisingeri* were collected from the mid to upper part of the Kralodvorian Králův Dvůr Formation, being particularly abundant in the so called Perník Bed (pelocarbonate), located in the uppermost part of the unit.

Portugal. – Portuguese specimens of *Vysocania* came from three different regions: from the Mós locality in Torre de Moncorvo Municipality of the Bragança District, in the northeastern part of the country (locality 2, Fig. 1A); from the Louredo locality in the Buçaco region of the Aveiro District, central Portugal (locality 3, Fig. 1A); and from the Pereiro locality in Mação Municipality of the Santarém District, central Portugal (locality 4, Fig. 1A). In the Torre

Series	Stage	Regional Stage	Simplified regional lithostratigraphy and stratigraphic distribution of <i>Vysocania</i> and <i>Octillaenus</i> spp.																																																		
			Czech Rep.			Portugal			Spain		Morocco																																										
Upper Ordovician	Sandbian	lower	Libeň Fm.	V. moravecii	V. cf. iberica	Chão do Amieiral Fm.	V. cf. iberica	Louredo Fm.	V. cf. iberica	Cabeço do Peão Fm.	Vysocania iberica	Fombuena Fm.	Vysocania sp.	Bancos Mixtos Fm.	V. iberica	Upper Ktaoua Fm.	Octillaenus maroccanus																																				
																		Berounian	middle	Vinice Fm.	V. panderi	V. cf. iberica	Louredo Fm.	V. cf. iberica	Cabeço do Peão Fm.	Vysocania iberica	Fombuena Fm.	Vysocania sp.	Bancos Mixtos Fm.	V. iberica																							
																															upper	Zahořany Fm.	V. panderi	V. cf. iberica	Louredo Fm.	V. cf. iberica	Cabeço do Peão Fm.	Vysocania iberica	Fombuena Fm.	Vysocania sp.	Bancos Mixtos Fm.	V. iberica											
																																											Bohdalec Fm.	V. panderi	V. cf. iberica	Louredo Fm.	V. cf. iberica	Cabeço do Peão Fm.	Vysocania iberica	Fombuena Fm.	Vysocania sp.	Bancos Mixtos Fm.	V. iberica

Figure 2. The chronostratigraphical, geographical and lithostratigraphical distribution of species of *Vysocania* Vaněk & Vokáč, 1997. Regional chronostratigraphic data after Bergström *et al.* (2009), Budil *et al.* (2011), Fatka *et al.* (2013) and Kraft *et al.* (2015).

de Moncorvo area, *V. cf. iberica* (Hammann, 1976) occurs in the lower part of the Chão do Amieiral Formation, of middle Berounian (global late Sandbian/early Katian) age (see Sá *et al.* 2005 for further information on the lithostratigraphy and palaeontology of the unit, and Pereira *et al.* 2015a for detailed stratigraphy of the locality and its trilobite assemblage). The base of this unit is marked by an oolitic ironstone bed (the Chôsavelha Bed), which correlates with the previously mentioned Zdice-Nučice Ironstone ‘Horizon’ of the Czech Republic. In the Buçaco region, *V. cf. iberica* occurs in oolitic ironstones of the Favaçal Bed at the base of the Louredo Formation. This bed correlates with the Chôsavelha Bed. Specimens of *V. iberica* from this region came from the overlying Porto de Santa Anna Formation, the base of which is also marked by an ironstone bed of late Berounian age, this species being the most common trilobite in it. In the Mação region, *V. iberica* occurs in the Cabeço do Peão Formation (see Young 1988 for lithostratigraphy, and Pereira *et al.* 2015b for information on the locality), which correlates with the Chão do Amieiral and Louredo formations.

Spain. – Spanish specimens of *Vysocania* came from two regions: the Fombuena locality in the province of Zaragoza, in the northeastern part of the country (locality 6, Fig. 1A); and the Almadén locality, in the Province of Ciudad Real in the south-central part of the country. The Fom-

buena specimens of *Vysocania* sp. came from the Huerva Member in the middle part of the Fombuena Formation (see Zamora *et al.* 2014 for information on local lithostratigraphy and the locality). This formation correlates with the middle Berounian units of Portugal. In the Ciudad Real region, *Vysocania iberica* occurs in the upper Berounian (ca. global lower Katian, Ka2) ‘upper coquinoid’ beds (type-horizon of this species) of the Bancos Mixtos Formation (see Hammann 1976 for fossil localities).

Morocco. – The specimens of *Octillaenus maroccanus* sp. nov. originate from Jbel Bou Ingarf (Fig. 1C), located between Zagora and Alnif, in the Drâa-Tafilalet region of southeastern Morocco. The type material was collected from about 40 m below the so-called ‘Ouzregui Bed’ (= *Dreyfussina struvei* bed of previous authors), in the uppermost part of the Upper Ktaoua Formation (Fig. 2), upper Kralodvorian (global upper Katian, Ka4). For further information on the stratigraphy and locality see Loi *et al.* (2010).

Systematic palaeontology

Repository acronyms are as follows: CMP – Miguel Pires collection and CMG – Marques Guedes collection, Departamento de Geologia from the Faculdade de Ciências da

Universidade de Lisboa (Lisbon, Portugal); MG – Museu Geológico de Lisboa (Lisbon, Portugal); MGUTAD – Museu de Geologia Fernando Real from the Universidade de Trás-os-Montes (Vila Real, Portugal); MGM – Museo Geominero of the Instituto Geológico y Minero de España (Madrid, Spain); MPZ – Museo de Ciencias Naturales de la Universidad de Zaragoza (Zaragoza, Spain); NHM – Natural History Museum (London, England); SMF – Senckenberg Museum (Frankfurt, Germany); NMP – Národní muzeum (Prague, Czech Republic); CGS – Česká geologická služba (Prague, Czech Republic) and MBHR – Muzeum Dr. Bohuslava Horáka v Rokycanech (Rokycany, Czech Republic).

Order Corynexochida Kobayashi, 1935
Suborder Illaenina Jaanusson, 1959
Family Illaenidae Hawle & Corda, 1847

Genus *Vysocania* Vaněk & Vokáč, 1997

Type species. – *Zbirovia vaneki* Šnajdr, 1958, Bohdalec Formation, upper Berounian (ca. global Ka2 stage slice, lower Katian; after Fatka *et al.* 2013 and Gutiérrez-Marco *et al.* 2017), Czech Republic.

Diagnosis. – Cranidium semi-circular in outline, very convex (sag. and tr.), weakly overhanging anterior margin; corset-shaped glabella (45–50% cranidial width posteriorly); axial furrows about one-third of the cranidial length; palpebral lobes of small to moderate size, located about their own length from posterior margin of cranidium; librigena subtrapezoidal, longer (exsag.) than wide (tr.), usually with pointed genal angle; thorax of nine segments, with narrow axis (approx. 40% of thoracic width anteriorly); rostral plate subtrapezoidal, with a length/width ratio about 28%, oblique and straight connective sutures (diverging forwards at about 70° to sagittal line) and a transverse depression posteromedially; hypostome width across anterior wings about 160% of sagittal length, lateral and posterior borders narrow and subparabolic in outline (medial part of posterior margin transverse), without shoulders, anterior lobe of the middle body strongly convex (tr. and sag.), occupying about 70% of the hypostomal length, posterior lobe slightly inflated, no maculae; pygidium subpentagonal to subelliptical in outline, length/width ratio from 55% to 70%; axis narrow (25% of pygidial width anteriorly), defined mainly by deflection of anterior margin; pygidial doublure with a medial forwardly directed deflection, about 35–50% of pygidial length.

Other species. – *Vysocania iberica* (Hammann, 1976), upper part of the Bancos Mixtos Formation, upper Berounian (ca. Ka2, lower Katian), Corral de Calatrava and Almadén areas, Spain; also known in the Queixoperra Member,

Cabeço do Peão Formation, middle Berounian (ca. upper Sandbian to lower Katian), Mação and in the lower half part of the Porto de Santa Ana Formation, upper Berounian (ca. Ka2, lower Katian), Buçaco, Portugal; *V. moraveci* sp. nov., Libeň Formation (also Letná Formation), lower Berounian (ca. Sandbian), Czech Republic; *V. panderi* (Barande, 1852), Zahořany Formation (also Vinice Formation), middle Berounian (upper Sandbian to lower Katian), Czech Republic.

Other references to Vysocania. – In Ibero-Armorica (Portugal, Spain and France), the records of *Vysocania* were previously reported as *Dysplanus* (*Zetillaenus*) or *Zetillaenus*. Although several authors reported the presence of *Vysocania* specimens in the Spanish middle Berounian, they have never been the subject of a systematic study: Kolb (1978, pl. 1, figs 11, 12) identified and figured *Dysplanus* (*Zetillaenus*)? sp. from the Piedra del Tormo and Huerva members of the Fombuena Formation from the Zaragoza region; Hammann *et al.* (1982) identified *Dysplanus* (*Z.*) cf. *ibericus* in the ‘Cuarцитas de Colmenarejos’ Formation from the Sierra del Tremedal region; Rábano (1984) mentioned the presence of *D. (Z.) ibericus* in the Collado de la Plata area, from a lateral equivalent of the Fombuena Formation; and Portero García *et al.* (1988) identified *Dysplanus* (*Zetillaenus*) sp. in the basal ironstone of the middle Berounian ‘Pizarras Cantera’ Formation from the Ciudad Real region. The material from the Huerva Member of the Fombuena Formation was revised for this work and assigned to *Vysocania* sp. Also, the identifications of *Dysplanus* (*Zetillaenus*) sp. in the upper Berounian to Kralodvorian Rosan Formation, from the Finistère region, France (Babin & Darboux 1982, p. 7), and *Stenopareia* sp. in the Lower Ktaoua Formation of Morocco (Álvarez *et al.* 2007, p. 32), probably correspond to *Vysocania*. Concerning the identifications of specimens from the Upper Ordovician Monte Orri Formation of Sardinia (*Zetillaenus* cf. *ibericus*; Hammann & Leone 1997, pl. 15, figs 1, 2, 5) and the Sart-Bernard Formation of Belgium (*Stenopareia* aff. *panderi*; Owens & Servais 2007, fig. 6e), it is not possible to confirm the generic assignments due to the rare occurrence, flattening and fragmentary preservation of the specimens coming from both of those units.

Remarks. – Confusion regarding the generic affinities of certain species here assigned to *Vysocania* dates back to the proposal of *Stenopareia* by Holm (1886) as a subgenus of *Illaeus* to encompass several species from the Ordovician of the Baltic region and the United Kingdom, as well as ‘*Illaeus*’ *panderi* from Bohemia. Jaanusson (1954) retained ‘*I.*’ *panderi* in *Stenopareia*, as did Šnajdr (1957) who also included the Czech species ‘*Illaeus*’ *oblitus* Barande, 1872 while highlighting differences between ‘*I.*’ *panderi* and Baltic *Stenopareia* species in the glabellar

muscle scars, rostral plate and hypostome. Most later authors (e.g. Shaw 2000, Bruthansová 2003) accepted the assignment of '*I.* *panderi*' to *Stenopareia*, but Vaněk & Valíček (2001) assigned '*I.* *panderi*' with question to *Vysocania*, which Bruthansová (2003) subsequently regarded as a junior synonym of *Stenopareia*. However, several dorsal and ventral morphological characters separate *Stenopareia* from *Vysocania*. *Stenopareia* has a relatively larger (tr.) glabella, usually defined by straighter dorsal furrows. In the librigenae, the genal angles are broadly rounded (in *Vysocania* they typically bear a short genal spine or show a pointed outline, except in *V. panderi* which has a slightly rounded genal angle). The glabellar musculature in *Stenopareia* (see Owen & Bruton 1980, pl. 3, fig. 10 and Bruton & Owen 1988, fig. 9n–p) is composed of four pairs of branched irregular scars, which decrease in size anteriorly (in *Vysocania* they are subcircular, according to Šnajdr 1957, text-fig. 3). The thorax of *Stenopareia* has a wider axis, corresponding to almost 50% of the anterior thoracic width, gently tapering posteriorly to 80% of its anterior width (the axis of *Vysocania* represents only 35% of the thoracic width anteriorly and is approximately 60% as wide posteriorly as anteriorly). The pygidial axis of *Stenopareia* is also wider, about 40% of the maximum pygidial width anteriorly (in *Vysocania* this ratio equals 25%). Therefore, the axial lobe of the entire exoskeleton of *Stenopareia* is wider (tr.) than in *Vysocania*.

The ventral morphology also shows noteworthy differences between representatives of both genera. *Stenopareia* displays a rostral plate with less oblique connective sutures (diverging 50° to sagittal line, compared to 70° in *Vysocania*; compare Fig. 7H herein and Owen & Bruton 1980, pl. 4, figs 4, 13). The posterior medial depression of *Vysocania* (Figs 4C, 7H–I) is absent in *Stenopareia*, in which the rostral plate has a dorsally and forwardly deflected subtriangular flange (Owen & Bruton 1980, pl. 4, fig. 9). The hypostome of *Stenopareia* has a different outline (compare Figs 7J, 8A and Owen & Bruton 1980, pl. 4, fig. 11), being subquadrangular, with rounded posterolateral corners (the *Vysocania* hypostome is roughly parabolic in outline, with the posterior margin transverse medially); the lateral border furrow is broad, the middle furrow is situated towards the back of the middle body, and the maculae are prominent (in *Vysocania* the middle furrow is oblique and located far forward, and there are no maculae).

Another important factor contributing to misunderstanding of the affinities of species belonging to *Vysocania* relates to *Zetillaenus* Šnajdr, 1957 (Fig. 11F–M), the type species of which is '*Illaenus*' *wahlenbergianus* Barrande, 1852 from the Králův Dvůr Formation of the Prague Basin. In erecting '*I.* *wahlenbergianus*', Barrande (1852, p. 683) contrasted the presence of a genal spine in this species with

what he described as the rounded genal angle of '*I.* *panderi*'. Consequently, the presence of a genal spine in *Vysocania iberica* may have played a role in relating that species to *Zetillaenus* (e.g. Thadeu 1947, p. 222; Hammann 1976, 1992; Young 1985; Romão *et al.* 1995; Hammann & Leone 1997). Although sharing similar general morphology, *Vysocania* and *Zetillaenus* clearly differ in a series of dorsal and ventral features. In *Zetillaenus* (Fig. 11F, K) the cephalon and pygidium are parabolic in outline (semicircular in *Vysocania*); and the longitudinal cranidial profile describes an arc of circle (the anterior cranidial region is more strongly convex in *Vysocania*). Concerning ventral features, *Zetillaenus* has a longer (sag.) rostral plate (Fig. 11G, I), the hypostome has distinct shoulders (Fig. 11H), (shoulders absent in *Vysocania*; Figs 7J, 8A); and the pygidial doublure is shorter (sag.; compare Fig. 8K and Fig. 11K).

Occurrence. – Upper Ordovician (Sandbian to lower Katian, Ka2; regional lower Berounian to upper Berounian stage) of Czech Republic, Portugal, Spain and possibly France.

Vysocania vaneki (Šnajdr, 1958)

Figure 3

- 1958 *Zbirovia vaneki* sp. nov.; Šnajdr, pp. 207–211, text-figs. 1, 2; pl. 1, figs 1–9; pl. 2, figs 10–20.
- 1966 *Stenopareia vaneki*. – Havlíček & Vaněk, p. 58, pl. 12, fig. 10.
- 1992 *Ulugtella vaneki*. – Hammann, p. 76.
- 1997 *Vysocania vaneki*. – Vaněk & Vokáč, p. 28, pl. 2, figs 12–21.
- 2001 *Vysocania vaneki*. – Vaněk & Valíček, p. 33.
- partim* 2003 *Stenopareia panderi*. – Bruthansová, pp. 178–181.
- 2003 *Stenopareia vaneki*. – Bruthansová, pp. 180, 181, figs 12f–i.

Holotype. – Internal mould of a pygidium (NMP L17619, figured by Šnajdr 1958, pl. 2, fig. 16 and Bruthansová 2003, fig. 12i).

Type horizon and locality. – Bohdalec Formation, upper Berounian (ca. Ka2, lower Katian), Hloubětín, Czech Republic.

Material. – Two enrolled specimens (CGS JV2294, JV2295-3); one cephalon (CGS JV2295-1); three cranidia (CGS JV10392, JV2295-3, MS3); one rostral plate (CGS Mu7); one thoracopygon (CGS JV2295-4); seven pygidia (CGS JV46, JV 1373, JV1431, JV2295-2, JV 10392, JV10403, JV 10421); three meraspides (cranidium, NMP L33192; pygidia, degree 4, NMP VH5176-1; degree 6, NMP A3).

Diagnosis. – Cranidium subrectangular; glabella narrow (width at lunette about 30% of maximum cranial width); axial furrows strongly bent at lunette; anterior section of facial suture straight; librigena with short genal spine; pygidium subpentagonal, length/width ratio about 65%; maximum pygidial width level with posterior end of axis; axis parabolic, anteriorly about 28–30% of maximum pygidial width, flat, axial furrows indistinct; doublure corresponding to 35% of sagittal pygidial length.

Remarks. – For detailed description see Šnajdr (1958, pp. 208–210) and Vaněk & Vokáč (1997, p. 28). The species *V. vaneki*, originally described as lacking eyes, was firstly established as a member of *Zbirovia* (Šnajdr 1958) and later tentatively re-assigned by Hammann (1992) to *Ulugtella* Petrunina in Repina *et al.* (1975). The revision of Vaněk & Vokáč (1997) showed the presence of palpebral lobes and a small genal spine in *V. vaneki* and these authors considered that all the *Vysocania* material from the upper Berounian Bohdalec Formation belongs to *V. vaneki*. However, Bruthansová (2003) considered both *Vysocania panderi* and *V. vaneki* to be represented in the Bohdalec Formation. Herein we follow the opinion of Vaněk & Vokáč (1997) that specimens from the Bohdalec Formation previously identified as *V. panderi* belong to *V. vaneki*. The latter species is differentiated from *V. panderi* by a narrower (tr.) glabella, a librigena with short genal spine, a shorter (sag.) pygidium with more subpentagonal outline, a shorter (sag.) pygidial doublure and the generally smaller size of the specimens.

Vysocania iberica (Hammann, 1976)

Figures 4, 5A–J

- 1908 *Illaenus* cf. *panderi* Barr. – Delgado, pp. 31, 57.
 1947 *Illaenus wahlenbergianus* Barrande. – Thadeu, p. 222, pl. 2, fig. 4.
partim 1976 *Dysplanus (Zetillaenus) ibericus* n. sp.; Hammann, pp. 51–53; text-fig. 5; pl. 2, figs 23–25; pl. 3, figs 26, 27–33, 35–37; non pl. 3, figs 28, 34, ? = *Cekovia loredensis* (Thadeu, 1947).
 1985 *Dysplanus (Zetillaenus)* sp. A. – Young, pp. 399, 400, pl. 48, fig. 15.

- 1985 *Dysplanus (Zetillaenus) ibericus* Hammann, 1976. – Young, p. 401; pl. 48, fig. 13.
 1987 *Dysplanus (Z.) ibericus*. – Gutiérrez-Marco & Rábano, p. 651.
 1995 *Zetillaenus* sp. – Romão *et al.*, p. 124.
 2000 *Zetillaenus* sp. – Romão, tab. 2.4.
 2011 *Zetillaenus wahlengergianus?* – Sá *et al.*, p. 31.

Holotype. – Internal mould of a cranidium (SMF 24860, figured by Hammann 1976, pl. 3, fig. 26).

Type horizon and locality. – Upper part ('upper coquinoid' beds) of the Bancos Mixtos Formation; locality Al I (in Hammann 1976, p. 37), north of Almáden, Spain.

Material. – Bancos Mixtos Formation (Spain): the type-material listed by Hammann (1976, p. 51).

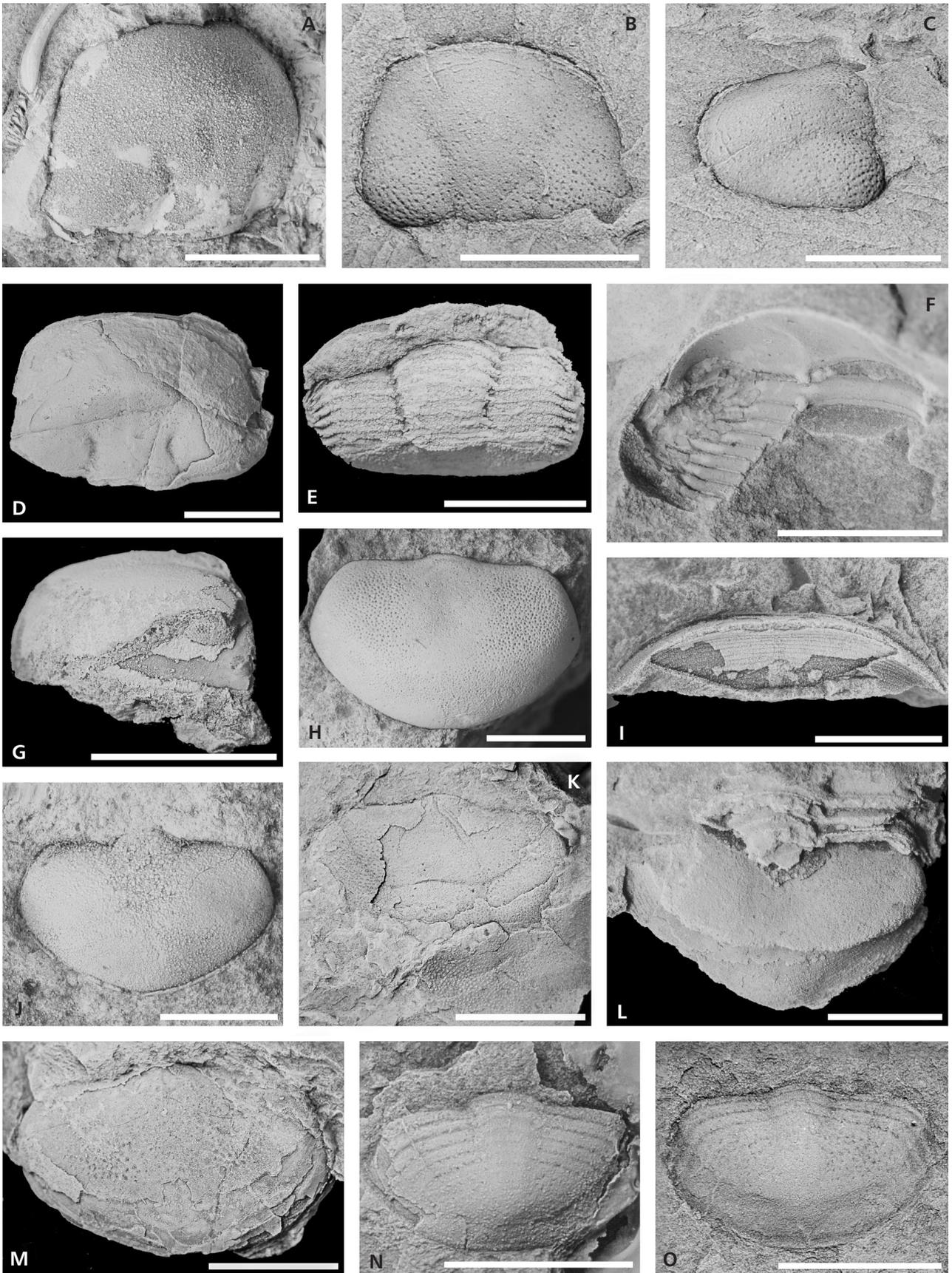
Queixoperra Member, Cabeço do Peão Formation (Portugal): thirteen exoskeletons (MG 91-2/100; CMG 0009, 0012a with counterpart 0012b, 0036, 0073; CMP 0095, 0142, 0200, 0295, 0297-1a with counterpart 0297-1b, 0297-2a with counterpart 0297-2b, 0300-1, 0300-2); nine cranidia (MGUTAD 15544, 15540, 15152, 15046, 15048, 15043, 15403; CMP0068, 0300-3); one pygidium (MGUTAD 15393); one meraspid pygidium (degree 7: CMP 0298).

Porto de Santa Anna Formation (Portugal): five exoskeletons (MG 2815-1, 214-a with counterpart 214-b, 214-2, 216, 216-2); four cranidia (MG 2813-8, 2813-11, 2813-12, 2805-2); one thoracopygon (MG 2815-2); four pygidia (MG2815-2, 2818-1, 2820-2, 2820-4); one meraspid exoskeleton (degree 8: MG 2805).

Diagnosis. – Cephalic axial furrows gently curved; palpebral lobes corresponding to 18–19% of cranial length; librigena bearing a short genal spine; pygidial length/width ratio from 58% to 65%; maximum pygidial width across posterior limit of axis; axial furrows indistinct; axis about 40% of pygidial length.

Description. – Cephalon semicircular, with maximum width (at posterior margin) of 160% of sagittal length. Cranidium strongly vaulted (sag. and tr.). Glabella corset-shaped, slightly exceeding height of fixigena in lateral

Figure 3. *Vysocania vaneki* (Šnajdr, 1958) from Bohdalec Formation, upper Berounian (ca. lower Katian, Ka2), Czech Republic; A – CGS MS3, oblique view (internal mould of a cranidium); B, C – NMP L33192, dorsal (B) and lateral (C) views (internal mould of a meraspid cranidium); D – CGS JV10392, dorsal view (internal mould of a cranidium); E – CGS JV2295e, dorsal view (internal mould of an enrolled exoskeleton); F – CGS JV2294a, oblique view (external mould of an enrolled specimen); G – CGS JV2295c (lateral view of a cephalon); H – CGS JV46, dorsal view (internal mould of a pygidium); I – CGS Mu7, ventral view (composite mould of cephalic doublure); J – CGS JV1431, dorsal view (internal mould of a pygidium); K – CGS JV10373b, dorsal view (internal mould of a pygidium and external mould of a cranial fragment); L – CGS Mu2, dorsal view (internal mould of a thoracopygon); M – CGS JV10421, dorsal view (internal mould of a pygidium); N – NMP VH5176-1, dorsal view (internal mould of a meraspid pygidium, degree 4); O – NMP A3, dorsal view (internal mould of a meraspid pygidium, degree 6). Scale bars: A, D–M = 5 mm; B, C, N, O = 2 mm.



view, merging with fixigenae anteriorly; maximum glabellar elevation at posterior quarter of glabellar length, sloping downwards anteriorly with strongly rounded profile; posterior glabellar width corresponding to 40–45% of maximum cranial width. Axial furrows moderately deep, broad (tr.), reaching from 35% to 40% of sagittal cranial length, converging gently forwards as far as lunette (opposite palpebral lobe), thereafter diverging, considerably shorter in front of lunette than behind, dying out abruptly. A broad (sag.) furrow (most distinct on internal moulds) close to the posterior edge of the glabella and parallel with it is interpreted as the occipital furrow. A median glabellar tubercle, only observed on the internal moulds, is present in front of the occipital furrow (Fig. 4A). Fixigenae subrectangular, width at palpebral lobes of 20% of maximum cephalic width; posterior border furrow on fixigena anteriorly arched, in transversal line with the occipital furrow. Posterior branch of facial suture short (exsag.), converging posteriorly; anterior branch of the facial suture divergent, initially straight in front of palpebral lobe, curving inwards anterior to 60% of cranial length. Palpebral lobe situated about its own length from posterior edge of cephalon. Librigenae subtriangular to subtrapezoidal, small, with a narrow and short (about 50% of length of eye) genal spine directed posterolaterally, its outer margin almost continuous in outline with the librigenal lateral margin; maximum librigenal width across posterior border. Doublure extending to the librigenal adaxial margin, bearing at least 15 equally spaced terrace ridges subparallel to the librigenal margin. Cranial sculpture of uniformly sized pits, anteriorly interspersed with slightly sinuous but rarely preserved terrace ridges. Rostral plate lenticular in outline, about 28% as wide at posterior end of connective suture as at anterior end, crossed posteromedially by shallow depression; connective suture converging backwards at 70° to sagittal line.

Thorax composed of nine segments, first three longer (sag., exsag.) than remainder, all of uniform width (tr.). Axis moderately arched (tr.), narrowing backwards, more strongly in last five segments, approximately 60% as wide posteriorly as anteriorly. Axial furrow deeper than on cephalon; subcircular axial processes visible on internal moulds. Pleurae successively increasing in width (tr.) posteriorly and fulcrum located successively farther from axial

furrow, at 30% of pleural width on first segment and 50% on last segment. Inner portion of pleura flat and smooth; outer portion deflected backwards and strongly downwards, bearing a broad (exsag., tr.) and smooth facet. Sculpture of indistinct pits on thoracic axis and inner part of pleurae, weaker than on cephalon and visible only on external surface; outer part of pleurae with subtransverse terrace ridges, about 20 on articulating facets.

Pygidial axis subtriangular to subparabolic, weakly convex, flat or even concave in transverse profile, anterior width about 25% of maximum pygidial width (level with posterior end of axis). Axial furrow not defined. Pleural areas smooth; pleural facet wide, about 50% of anterior pleural width, oriented at about 110° to sagittal line, with terrace ridges with similar configuration as on thoracic pleural facets. Dorsal surface of pygidium, excluding pleural facets, with similar pitted sculpture as cranium, less dense in axial than pleural region. Doublure corresponding to 50% of pygidial length medially, decreasing to 40% laterally; inner margin with forward deflection medially; sculpture of about 15 similarly spaced (about 0.25 mm) terrace ridges running subparallel to pygidial margin, becoming more sinuous medially.

Remarks. – The type-material of *Vysocania iberica*, deposited in the Senckenberg Museum (Frankfurt, Germany), comprising deformed and fragmentary specimens, was revised for this work. We consider that among this type-material at least two different illaenids are represented: *V. iberica* and *Cekovia* cf. *loredensis* (see synonymy of *Vysocania iberica* herein).

Vysocania cf. *iberica* (Hammann, 1976)

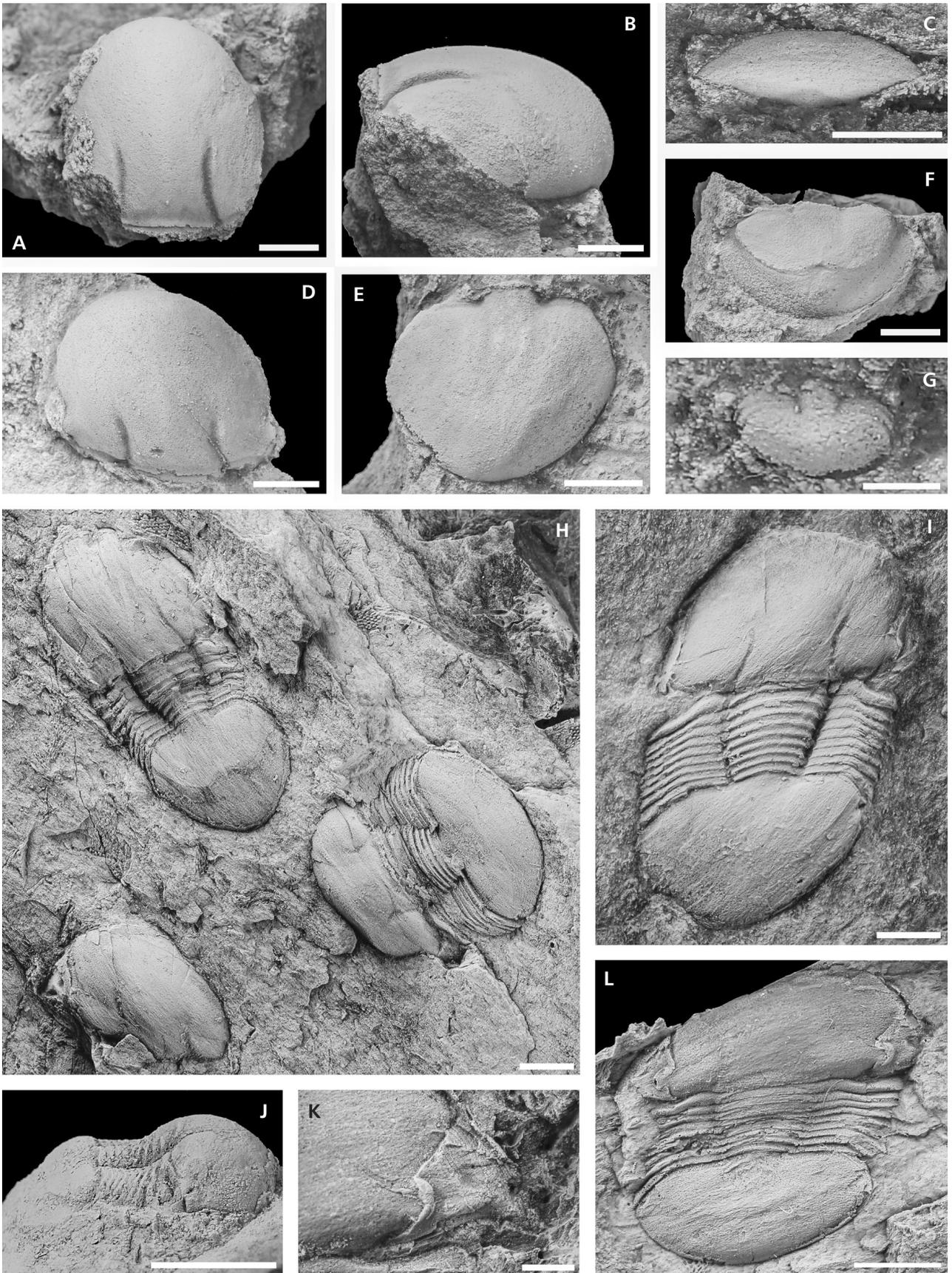
Figure 5K–M

2015a *Stenopareia* (*Vysocania*) sp. – Pereira *et al.*, p. 62, figs 1c, d.

Material. – Chão do Amieiral Formation (Torre de Moncorvo, Portugal): one cranium (MGUTAD 15703); one pygidium (MGUTAD 15704).

Louredo Formation (Buçaco, Portugal): one pygidium as counterpart moulds (MGUTAD 15705a, 15705b).

Figure 4. A–G – *Vysocania iberica* (Hammann, 1976) from Bancos Mixtos Formation, upper Berounian (*ca.* lower Katian, Ka2), Almáden, Spain; A, B – SMF 24873 (paratype), dorsal (A) and lateral (B) views (internal mould of a cranium); C – SMF 24872 (paratype), ventral view (internal mould of a rostral plate); D – SMF 24873-2 (paratype), dorsal view (internal mould of a cranium); E – SMF 24874 (paratype), dorsal view (internal mould of a pygidium); F – SMF 24871 (paratype), dorsal view (composite mould of a pygidium); G – SMF 24861 (paratype), dorsal view (internal mould of a meraspid pygidium, degree 7). • H–L – *Vysocania iberica* (Hammann, 1976) from Cabeço do Peão Formation, middle Berounian (upper Sandbian/lower Katian), Mação, Portugal. • H – CMP 0300, dorsal view (internal moulds of two exoskeletons and one cephalon); I – CMP 0142, dorsal view (internal mould of an exoskeleton); J – CMP 0200, lateral view (internal mould of an exoskeleton); • K, L – CMG 0012, dorsal views. Detail of the librigena of the internal mould of an exoskeleton (K), and latex cast of the external mould of the exoskeleton (L). Scale bars = 5 mm; except in G, K (= 2 mm).



Remarks. – The specimens are too poorly preserved for specific identification. However, the presence of the main diagnostic characters of *Vysocania* supports the generic assignment. One cranidium (Fig. 5K) preserves a gently curved axial furrow typical of *V. iberica*. Based on this feature, and considering the geographical and chronostratigraphic context, this material is tentatively assigned to the only known Iberian species of *Vysocania*.

Occurrence. – Chão do Amieiral Formation, Torre de Moncorvo (Mós), and Louredo Formation, Buçaco (Louredo), from the middle Berounian (*ca.* upper Sandbian to lower Katian) of Portugal.

***Vysocania moraveci* sp. nov.**

Figure 6

- partim* 1957 *Stenopareia panderi*. – Šnajdr, pp. 160–166, pl. 7, fig. 3 [*non* pl. 7, figs 4–7 = *Vysocania panderi*].
2001 *Harrisia rumpalensis* Moravec (MS). – Vaněk & Valíček, p. 33 [*nomem nudum*].
2001 *Harrisia omega* Moravec (MS). – Vaněk & Valíček, p. 33 [*nomem nudum*].
2002 *Harrisia rumpalensis*. – Moravec, p. 56 [*nomen nudum*].
partim 2003 *Stenopareia panderi*. – Bruthansová, pp. 178–180.
2006 *Harrisia rumpalensis* Moravec (MS). – Vaněk & Valíček, p. 28 [*nomen nudum*].
2006 *Harrisia omega* Moravec (MS). – Vaněk & Valíček, p. 28 [*nomen nudum*].
2011 *Stenopareia panderi*. – Mergl & Budil, p. 111, fig. 1i, j.

Holotype. – Internal mould of a pygidium (CGS JV10511; Fig. 6A, B).

Paratypes. – Two incomplete exoskeletons (CGS JV10497, JV10505); seven cranidia (CGS JV10397-2, JV10400-2, JV10507, JV10510–JV10512, JV10513-1); fourteen pygidia (CGS JV10395, JV10397-1, JV10400-1, JV10504, JV10506, JV10507, JV10509, JV10513-2, JV10514–JV10519).

Type horizon and locality. – Řevnice Quartzite of the Libeň Formation (lower Berounian, *ca.* Sandbian), Zbiroh (49° 49' 55" N; 13° 47' 12" E), Plzeň region, Rokycany district, Czech Republic.

Other material. – Libeň Formation: one cranidium (CGS MS1746); four pygidia (CGS MS1750, MS1751, MS10658; MBHR 6166).

Letná Formation: one incomplete exoskeleton (NMP 23937-3); one cephalon (CGS JV 1502); two cranidia (NMP 23937-1, 23937-2).

Etymology. – After Josef Moravec (Nepomuk, Czech Republic), a private investigator in palaeontology. He was a pioneer in recognizing the different identity of *Vysocania* specimens from the Libeň Formation.

Diagnosis. – Exoskeleton small (10–15 mm of sagittal length), with moderately curved cephalic axial furrows extending almost 50% of cranidial length; librigena with pointed genal angle; pygidium with posterior margin slightly ogival, length/width ratio low (from 55% to 65%); maximum pygidial width across posterior limit of axis; axis about 40% of pygidial length and 27% of maximum pygidial width, parabolic, weakly convex, defined by faint axial furrows.

Description. – Cephalon semicircular, with maximum width at posterior margin. Cranidium strongly vaulted (sag. and tr.). Glabella corset-shaped, slightly exceeding height of fixigena in lateral view, merging with fixigenae anteriorly; maximum glabellar elevation at posterior quarter of glabellar length, sloping downwards anteriorly with strongly rounded profile; posterior glabellar width corresponding to 45–50% of maximum cranidial width. Axial furrows moderately deep, broad (tr.), converging gently forwards as far as lunette (opposite palpebral lobe), thereafter diverging. A broad (sag.) furrow (most distinct on internal moulds) close to the posterior edge of the glabella and parallel with it is interpreted as the occipital furrow. Fixigenae subrectangular; posterior border furrow on fixigena anteriorly arched, in transversal line with the occipital furrow. Posterior branch of facial suture short (exsag.), converging posteriorly; anterior branch of the facial suture divergent, initially straight in front of palpebral lobe, curving inwards at the anterior cephalic margin. Palpebral lobe situated about half of its own length from posterior edge of cephalon. Librigenae subtriangular to subtrapezoidal; maximum librigenal width across posterior border. Thorax composed of nine segments, all of uniform width (tr.). Axis moderately arched (tr.), narrowing backwards, more strongly in last five segments, approximately 55% as wide posteriorly as anteriorly. Axial furrow moderately deep; subcircular axial processes visible on internal moulds. Inner portion of pleura flat and smooth; outer portion deflected backwards and strongly downwards, bearing a broad (exsag., tr.) facet.

Pygidial axis subparabolic, weakly convex to flat in transverse profile. Pleural areas smooth; pleural facet wide, about 50% of anterior pleural width. Dorsal surface of pygidium with pitted sculpture as cranidium. Doublure corresponding to 50% of pygidial length medially, slightly decreasing laterally; inner margin with forward deflection medially.

Remarks. – In general morphology, *V. moraveci* is similar

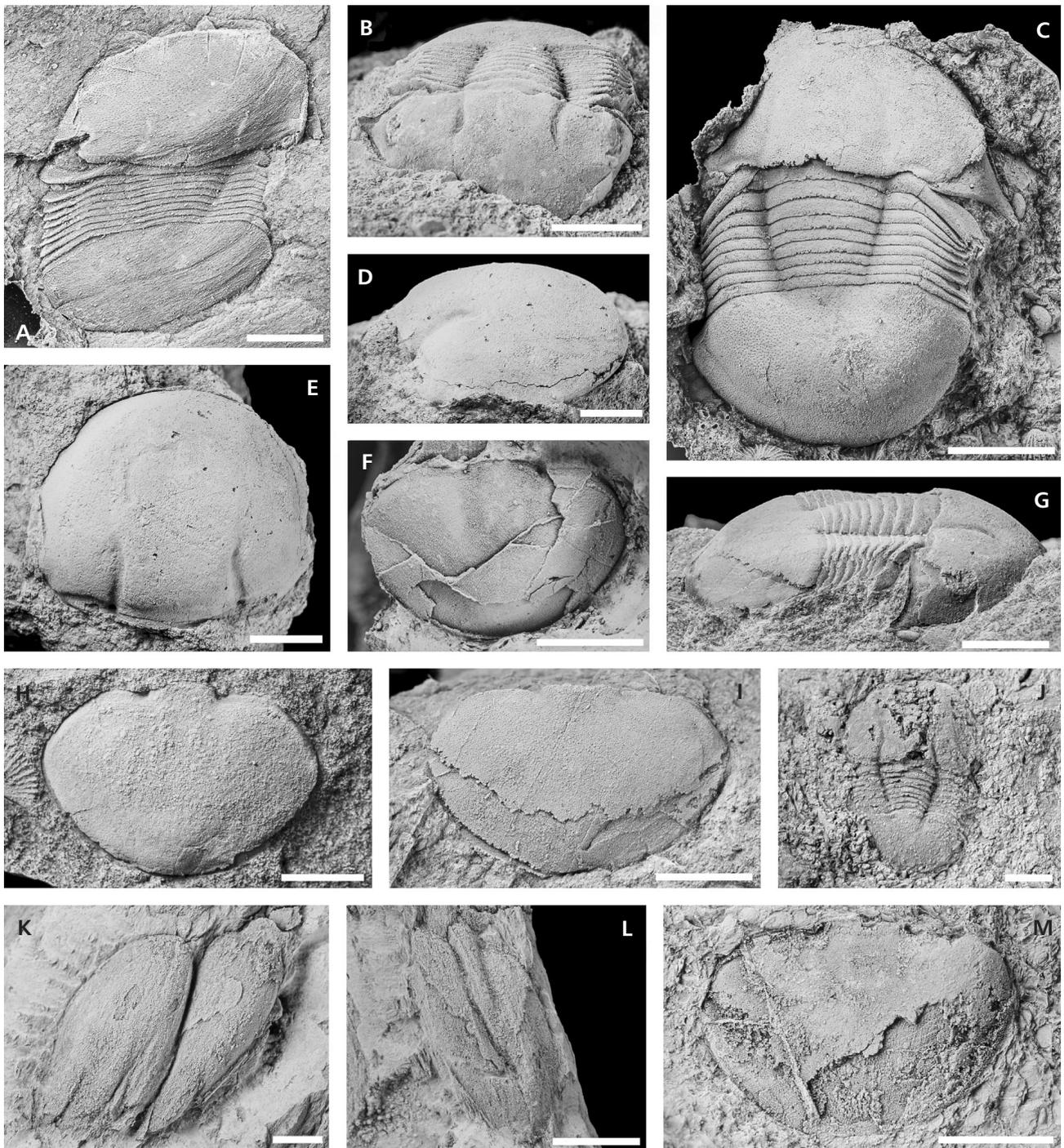


Figure 5. A–J – *Vysocania iberica* (Hamman, 1976) from Porto de Santa Anna Formation, upper Berounian (ca. lower Katian, Ka2), Buçaco, Portugal; A – CMP 0295, dorsal view (latex cast of the external mould of an exoskeleton); B, C, G – MG 214, anterior (B), dorsal (C) and lateral (G) views (internal moulds, B, G, and latex cast of the external mould, C, of an exoskeleton); D, E – MG 2813-11, lateral (D) and dorsal (E) views (internal mould of a cranidium); F – MG 2533, dorsal view (internal mould of a pygidium); H – MG 2818-1, dorsal view (internal mould of a pygidium); I – MG 2820-4, dorsal view (composite mould of a pygidium); J – MG 2805, dorsal view (internal mould of a meraspid exoskeleton, degree 8). • K–M – *Vysocania cf. iberica* (Hamman, 1976); K–L – from the Chão do Amieiral Formation, middle Berounian (upper Sandbian/lower Katian), Moncorvo, Portugal; M – from the Louredo Formation, middle Berounian (upper Sandbian/lower Katian), Buçaco, Portugal; K – MGUTAD 15703, dorsal view (internal mould of a cranidium); L – MGUTAD 15704, dorsal view (composite mould of a pygidium); M – MGUTAD 15705, dorsal view (composite mould of a pygidium). Scale bars = 5 mm except in J (= 2 mm).

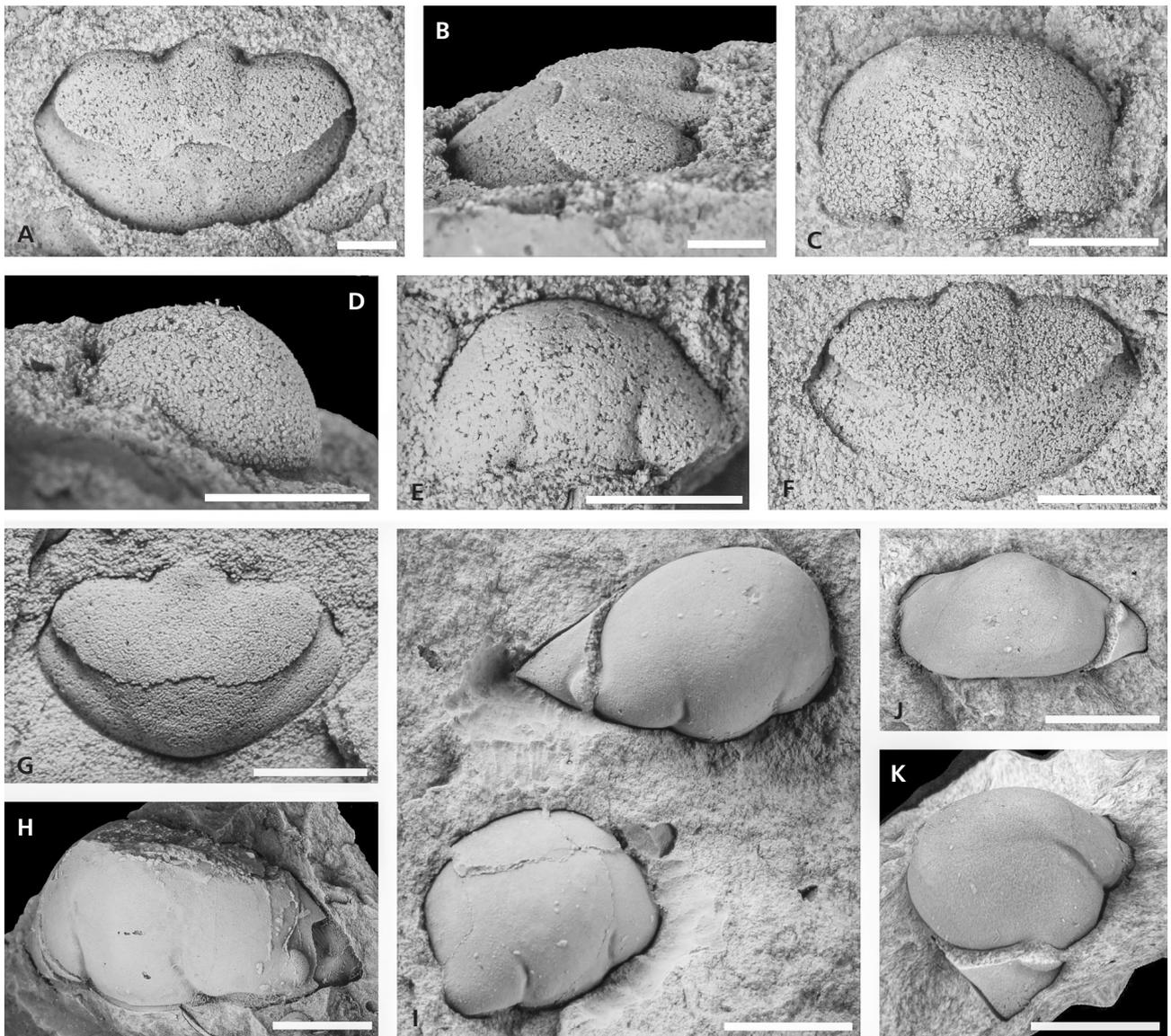
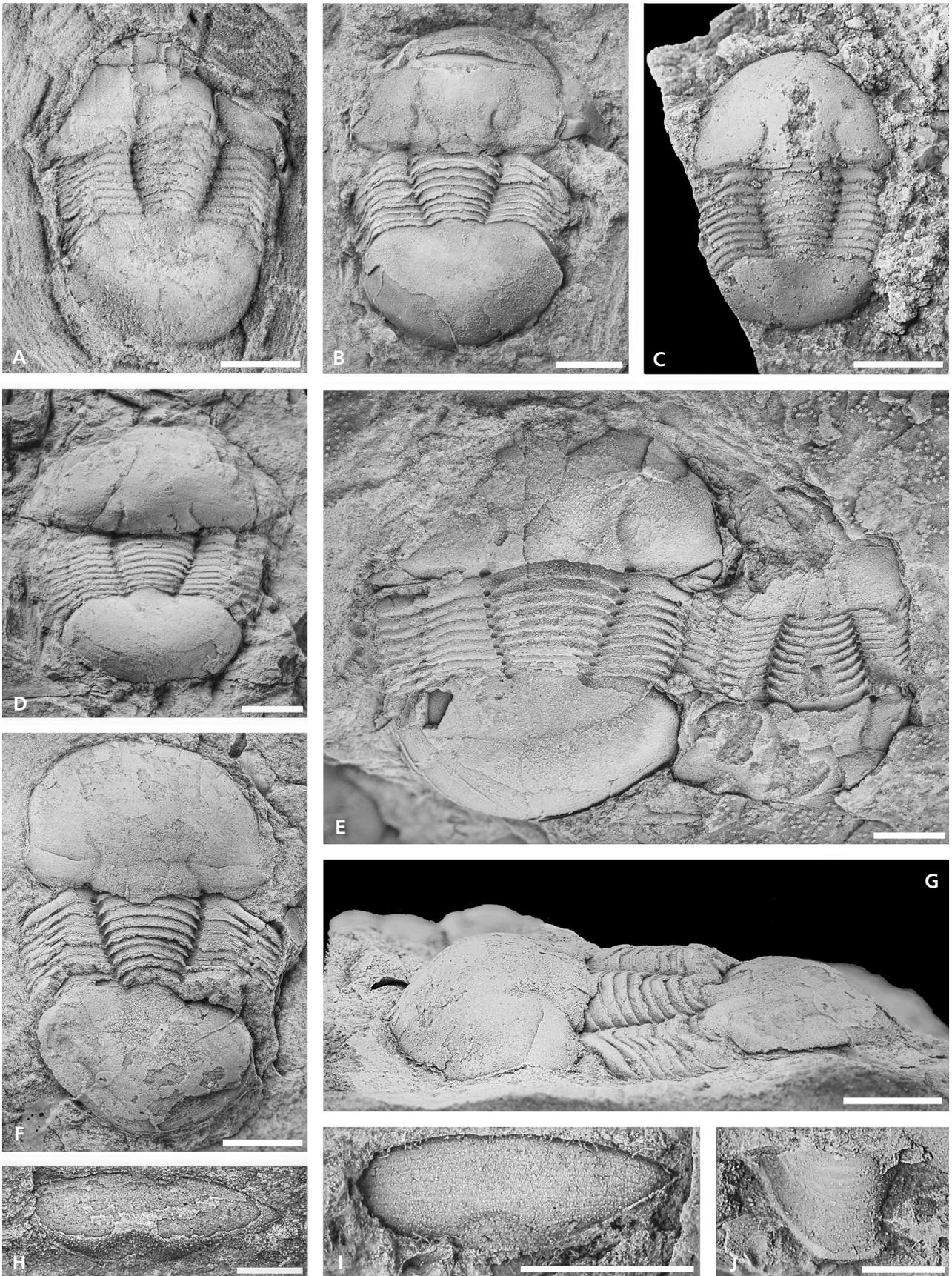


Figure 6. A–G – *Vysocania moraveci* sp. nov. from Libeň Formation, lower Berounian (ca. Sandbian), Zbiroh (A–F) and Rumpál (G), Czech Republic; A, B – CGS JV10511 (holotype), dorsal (A) and lateral (B) views (composite mould of a pygidium); C, D – CGS MS1746, dorsal (C) and lateral (D) views (internal mould of a cranium); E – JV 10512a, dorsal view (internal mould of a cranium); F – CGS MS10658, dorsal view (composite mould of a pygidium); G – MBHR 6166, dorsal view (composite mould of a pygidium). • H–K – *Vysocania moraveci* sp. nov. from Letná Formation, lower Berounian (ca. Sandbian), Chrustenice, Czech Republic; H – CGS JV1502, oblique view (internal mould of a cephalon); I, J, K – NMP 23937, dorsal (I), anterior (J) and lateral (K) views (internal moulds of one cephalon and one cranium). Scale bars = 5 mm except in A, B (=2 mm).

Figure 7. A – *Vysocania panderi* (Barrande, 1852) from Vinice Formation, middle Berounian (upper Sandbian/lower Katian), Trubín, Czech Republic. NMP 35050, dorsal view (internal mould of an exoskeleton). • B–J – *Vysocania panderi* (Barrande, 1852) from Zahořany Formation, middle Berounian (upper Sandbian/lower Katian), Czech Republic; B – CD 1112, dorsal view (internal mould of an exoskeleton); C – NMP L46553, dorsal view (internal mould of an exoskeleton); D – NMP L15610, dorsal view (internal mould of an exoskeleton); E – NHM42609-1, dorsal view (internal moulds of two exoskeletons); F, G – NHM 42609-2, dorsal (F) and lateral (G) views (internal mould of an exoskeleton); H – NMP L15605, ventral view (internal mould of a rostral plate); I – CGS JV10407, ventral view (internal mould of a rostral plate); J – JV 2297, ventral view (internal mould of a hypostome). Scale bars = 5 mm except in H–J (= 2 mm).



to other species of the genus. Previous records of *Vysocania* in the lower Berounian Libeň and Letná formations were assigned to *V. panderi* by Šnajdr (1957) and Bruthansová (2003). However, Jiří Vaněk, in a late stage of his career, considered that specimens from Libeň represent a new species, which he named '*Stenopareia vespertilio*' – *nomen nudum* – on the labels of his collection housed in the Česká geologická služba in Prague. This name was not included in the trilobite indexes published by Vaněk & Valíček (2001, 2006), but it certainly corresponds to the new genus and new species suggested by Moravec (2002), but never validly erected, *Harrisia rumpalensis* – *nomina nuda*. On the other hand, the specimens from the Letná Formation were named as *Harrisia omega* Moravec – *nomina nuda* – (Vaněk & Valíček 2001). In this work, we revised the collections of those authors and other *Vysocania* specimens from the Libeň and Letná formations and we consider all of the material to be conspecific, the only difference being the slightly larger size of specimens from the latter unit.

Vysocania moraveci differs from the stratigraphically younger *V. panderi* in having a pointed genal angle (Fig. 6I; in *V. panderi* it is rounded), lower pygidial length/width ratio (55–65% instead of 65–70% in *V. panderi*), ogival pygidial posterior margin (in *V. panderi* it is semicircular) and relatively longer pygidial axis (50% of pygidial length, instead of 35–40%).

Vysocania panderi (Barrande, 1852)

Figures 7, 8A–K

- partim* 1852 *Illaeus panderi*; Barrande, pp. 682–683, pl. 30, figs 4–9; pl. 35, figs 21–25 [non pl. 30, figs 10, 11 = *Octillaenus hisingeri*].
- partim* 1957 *Stenopareia panderi*. – Šnajdr, pp. 160–166, text-figs 3, 7; pl. 7, figs 4–7 [non pl. 7, fig. 3 = *V. moraveci* sp. nov.].
- 2001 *Vysocania* (?) *vanessa* Moravec (MS). – Vaněk & Valíček, p. 34 [nomen nudum].
- partim* 2003 *Stenopareia panderi*. – Bruthansová, pp. 178–180, fig. 11a–c.
For a complete synonymy see Šnajdr (1957, p. 160).

Lectotype. – Designated by Šnajdr (1957); internal mould of a complete exoskeleton (NMP L15607, figured by Barrande 1852, pl. 35, fig. 21 and Bruthansová 2003, fig. 11a).

Type horizon and locality. – Zahořany Formation, middle Berounian (ca. upper Sandbian/lower Katian), Zahořany, Czech Republic.

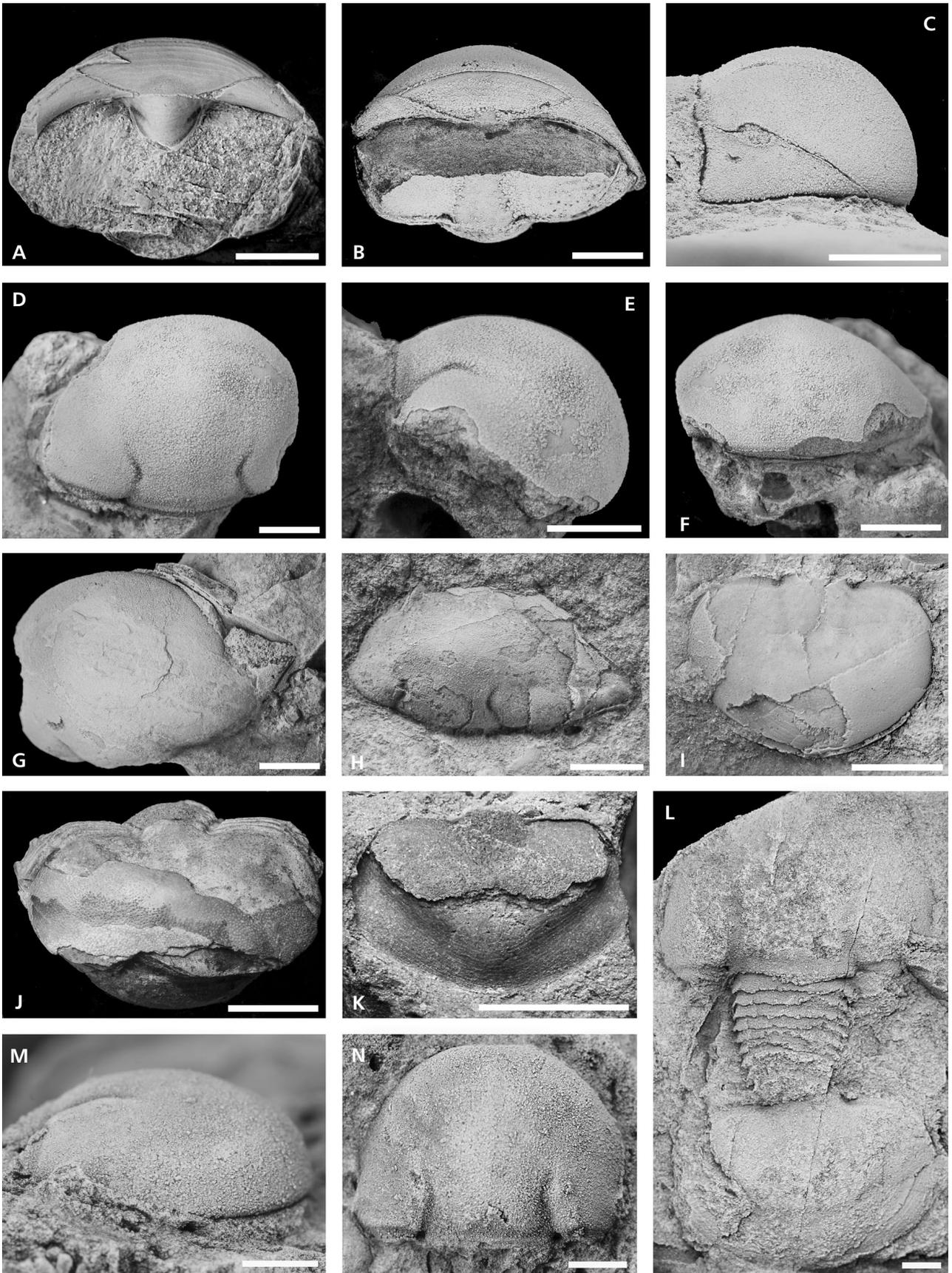
Material. – Vinice Formation: three exoskeletons (NMP 35050-6–35050-8).

Zahořany Formation: ten exoskeletons (NHM 42609-1, 42609-2; NMP L15610, 46553; CD1112, 6465, 35050-2, 35050-4, 35050-5; CGS JV1682); one enrolled exoskeleton (NHM 42458-1); seven cephalae (NHM42458-3; NMP 35050-3, L46555; CGS JV2288, JV2289, JV2735, JV10384); one cranidium (NMP 35050-1); one cephalothorax (CGS JV2737-2); five pygidia (NHM 42458-4; NMP L46556; CGS JV2737-1, JV10405, JV10428); three articulated rostral plate + hypostome (NMP L15605, L46554; CGS MS21); two rostral plates (NMP L15605; CGS JV10407); one isolated hypostome (CGS JV2297).

Diagnosis. – Cephalic axial furrows strongly curved, extending about $\frac{1}{3}$ of cranidial length; librigena with slightly rounded genal angle; pygidium with evenly curved posterior margin, length/width ratio high (from 65% to 70%); maximum pygidial width across posterior limit of axis; axis parabolic, about 35–40% of pygidial length and 25% of maximum pygidial width, flat to slightly concave, axial furrows indistinct; doublure long (55–60% of sagittal pygidial length), inner edge with a wide (tr.) forward deflection medially.

Remarks. – For detailed descriptions see Šnajdr (1957, pp. 160–166) and Bruthansová (2003; pp. 178, 179). *Vysocania panderi* differs from other species of the genus in having a broadly rounded anterior cranidial margin, a slightly rounded genal angle (the other species all have small genal spines or pointed genal angles) and longer (sag.) pygidium (higher length/width ratio). The pygidial length/width ratio varies from 0.65 to 0.70 (based on dozens of specimens). The 0.82 ratio indicated by Mergl & Budil (2011) in the lectotype of *V. panderi* (NMP L15607) is certainly due to deformation, to longitudinally stretched specimens.

Figure 8. A–K – *Vysocania panderi* (Barrande, 1852) from Zahořany Formation, middle Berounian (upper Sandbian/lower Katian), Czech Republic; A – NMP L46554, ventral view (internal mould of a cephalon with hypostome); B – NHM 42458-1, ventral and dorsal views (internal mould of an enrolled exoskeleton, detail of the cephalic doublure and the fragmentary pygidium); C – NHM 42458-3, lateral view (internal mould of a cephalon); D, E, F – NMP 35050, dorsal (D), lateral (E) and anterior (F) views (internal mould of a cranidium); G – CGS JV2289, oblique view (internal mould of a cephalon); H – NMP L46555, dorsal view (internal mould of a cephalon); I – NMP L46556, dorsal view (composite mould of a pygidium); J – NHM I3410, posterodorsal view of pygidium (internal mould of an enrolled exoskeleton); K – NHM It13240, dorsal view (composite mould of a pygidium). • L–N – *Vysocania* sp. from the Huerva Member, Fombuena Formation, middle Berounian (upper Sandbian/lower Katian), Zaragoza, Spain; L – MPZ 2016/121, dorsal view (internal mould of an exoskeleton); M, N – MPZ 2016/123, lateral (M) and dorsal (N) views (internal mould of a cranidium). All scale bars = 5 mm.



Specimens of *Vysocania* are rare in the middle Berounian Vinice Formation and were previously assigned to a new species by Moravec, who named it *Vysocania* (?) *vanessa* – *nomen nudum* – (Vaněk & Valíček 2001). However, the specimens from this formation agree in morphology with those of *V. panderi* from the overlying Zahořany Formation, and consequently are here assigned to that species. Specimens from the Libeň and Letná formations previously assigned to *V. panderi* (e.g. Barrande 1852, Šnajdr 1957, Bruthansová 2003) are here included in *V. moraveci*.

***Vysocania* sp.**

Figure 8L–N

1978 *Dysplanus* (*Zetillaenus*)? sp. – Kolb, pl. 1, figs 11, 12.

Material. – Fombuena Formation (Zaragoza, Spain): one exoskeleton (internal mould: MPZ 2016/121a; external mould: MPZ 2016/121b); two cranidia as counterpart moulds (MPZ 2016/122a, 2016/123a, 2016/122b, 2016/123b).

Remarks. – The specimens show the diagnostic features of *Vysocania iberica*, except that one dorsal exoskeleton with the right librigena attached seems to have a rounded genal angle (Fig. 8L, observed in the external mould). Additional material from the Fombuena Formation, not examined for this study, also lacks a genal spine (Samuel Zamora, personal communication 2016). Further sampling in this region may clarify the species assignment. The specimens from the Fombuena Formation are among the largest known for the genus, reaching 60 mm in length.

Occurrence. – Huerva Member, Fombuena Formation, Fombuena (see Zamora *et al.* 2014), middle Berounian (upper Sandbian/lower Katian), Zaragoza, Spain.

Genus *Octillaenus* Salter, 1867

Type species. – *Illaeus hisingeri* Barrande, 1846, Králův Dvůr Formation, Kralodvorian (ca. global Ka3–4 stage slices, upper Katian), Czech Republic.

Diagnosis. – Cranidium semi-circular in outline, very convex (sag. and tr.); glabella corset-shaped, comprising

45–50% of cranidial width posteriorly; axial furrows up to 60% of cranidial length; palpebral lobes of small to moderate size, located about half their own length from posterior margin of cranidium; librigena subtrapezoidal, longer (exsag.) than wide (tr.), with a short genal spine; thorax of eight segments, with narrow axis (approx. 40% of thoracic width anteriorly); rostral plate subtrapezoidal, with a length/width ratio about 30%, oblique and almost straight connective sutures (diverging about 70° to sagittal line) and a transverse posterior medial depression; pygidium subpentagonal to semicircular in outline, axis narrow (25% of pygidial width anteriorly), defined mainly by deflection of anterior margin; pygidial doublure about 30% of pygidial length, inner margin with a short (sag.), forwardly directed cusp medially.

Other species. – *Octillaenus maroccanus* sp. nov., Upper Ktaoua Formation, Kralodvorian (ca. Ka4, upper Katian), Bou Ingarf (30° 34' 22" N; 5° 34' 20" W), Zagora, Morocco.

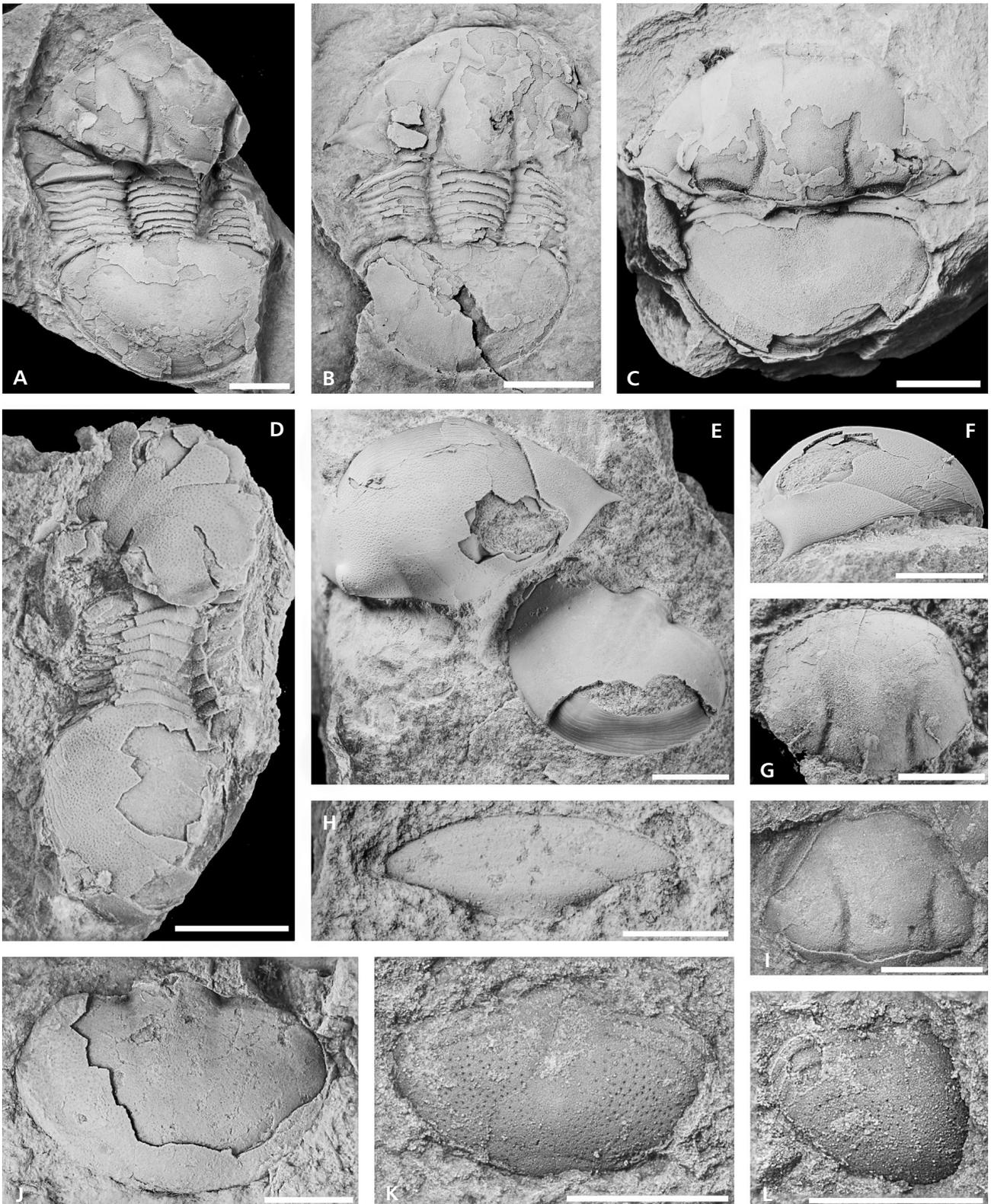
Occurrence. – Upper Ordovician (upper Katian, Ka3–4; regional Kralodvorian) of Czech Republic and Morocco. Moreover, the genus was previously documented in Sweden, as *Octillaenus hisingeri* by Troedsson (1924) and in the United Kingdom as *Octillaenus* aff. *hisingeri* by Ingham (1970; pl. 3, figs 15–25), but it is not possible to confirm the generic assignment on the basis of the published illustrations.

***Octillaenus hisingeri* (Barrande, 1846)**

Figure 9

- 1846 *Illaeus hisingeri*; Barrande, p. 14.
- 1852 *Illaeus hisingeri*. – Barrande, p. 681, pl. 29, figs 25–29.
- 1872 *Illaeus oblitus*. – Barrande, p. 73, pl. 15, figs 44, 45.
- 1957 *Octillaenus hisingeri*. – Šnajdr, pp. 155–159, pl. 7, figs 1, 2.
- 1957 *Stenopareia oblita*. – Šnajdr, pp. 166–169, pl. 7, figs 8–10.
- 1980 *Octillaenus hisingeri*. – Přibyl & Vaněk, pp. 270, 271, pl. 2, figs 3–5.
- 2000 *Octillaenus hisingeri*. – Shaw, p. 376, pl. 1, figs 3–5, 7, 14.
- 2000 *Stenopareia oblita*. – Shaw, pp. 376–377, pl. 1, figs 11–13, 15.

Figure 9. *Octillaenus hisingeri* (Barrande, 1846) from the Králův Dvůr Formation, Kralodvorian (upper Katian), Czech Republic; A – NMP L15163, dorsal view (internal mould of an exoskeleton); B – CGS JV1597a, dorsal view (internal mould of an exoskeleton); C – CGS JV1597b, dorsal view (internal mould of an exoskeleton with cephalon overlapping the thorax); D – CGS PB 2001, dorsal view (internal mould of an incomplete exoskeleton); E, F – NMP L32849, dorsal (E) and lateral (F) views (internal mould of a cephalon and composite mould of a pygidium); G – NMP 6-00, dorsal view



(internal mould of a cranidium); H – CGS JV9689, ventral view (internal mould of a rostral plate); I – CGS JV2743b-2, dorsal view (internal mould of a meraspis cranidium); J – CGS JV1641, dorsal view (internal mould of a pygidium); K – CGS JV2743a, dorsal view (internal mould of a meraspis pygidium, degree 6); L – CGS JV2743b-1, dorsal view (internal mould of a meraspis pygidium, degree 6?). Scale bar in A–G, J = 5 mm; in H–I, K–L = 2 mm.

- 2003 *Octillaenus hisingeri*. – Bruthansová, pp. 175, 176, figs 8a, b, d.
 2003 *Stenopareia oblita*. – Bruthansová, p. 180, fig. 12a–e.
 2009 *Stenopareia oblita*. – Fatka & Mergl, p. 88, fig. 12c.
 For a complete synonymy list see Šnajdr (1957, pp. 155, 156, 166) and Bruthansová (2003; pp. 175, 176, 180).

Lectotype. – Designated by Šnajdr (1957, p. 156); dorsal exoskeleton (NMP L15170, figured by Barrande 1852, pl. 29, figs 27, 28; Příbyl & Vaněk 1980, pl. 2, fig. 3 and Shaw 2000, pl. 1, fig. 5).

Type horizon and locality. – Králův Dvůr Formation, Kralodvorian (ca. Ka4, upper Katian), Králův Dvůr, Beroun, Czech Republic.

Material. – Thirteen exoskeletons (CGS PB2001a, with counterpart 2001b, PB2002, L32844, L32845, L32846, L15163, L15169, L15173, L15163, L15169, L15171, L15173; CGS JV1597); one cephalon (NMP L32849); three cranidia (NMP 8-200, 6-00-1; CGS JV7588); one rostral plate (CGS JV9689); four pygidia (NMP 6-00-2, L32849; CGS JV2739-1, 2739-2); seven meraspides (cranidia: NMP A1-1, A1-2; CGS JV2743b-2; pygidium degree 7: NMP L33188, degree 6: CGS JV2743a, JV2743b-1, degree 5: CGS JV2743c).

Remarks. – *Illiaenus oblitus* was synonymized with *Octillaenus hisingeri* by Příbyl & Vaněk (1980), Vaněk & Voják (1997) and Vaněk & Valíček (2001, 2006), but Hammann (1992), Shaw (2000) and Bruthansová (2003) rejected this synonymy. After identifying *Stenopareia* cf. *oblita* in the Kralodvorian ‘Caliza de Cistoideos’ Formation from Spain, Hammann (1992) distinguished *O. hisingeri* and *‘I.’ oblitus* based on the characters of the Spanish material, which is certainly neither conspecific nor congeneric with the Czech type material of *‘I.’ oblitus*. The Spanish specimens, revised for this work, have smaller eyes, located further anteriorly, subtriangular librigenae with broadly rounded genal angle, very thick cuticle with coarser, deeper and denser pitting, more convex cranidial sagittal profile, longer pygidia with a strong pentagonal outline and the pygidial axis is segmented. For these reasons, we do not agree with Shaw (2000) that Hammann (1992) gave clear reasons why Příbyl & Vaněk’s (1980) attempt to synonym-

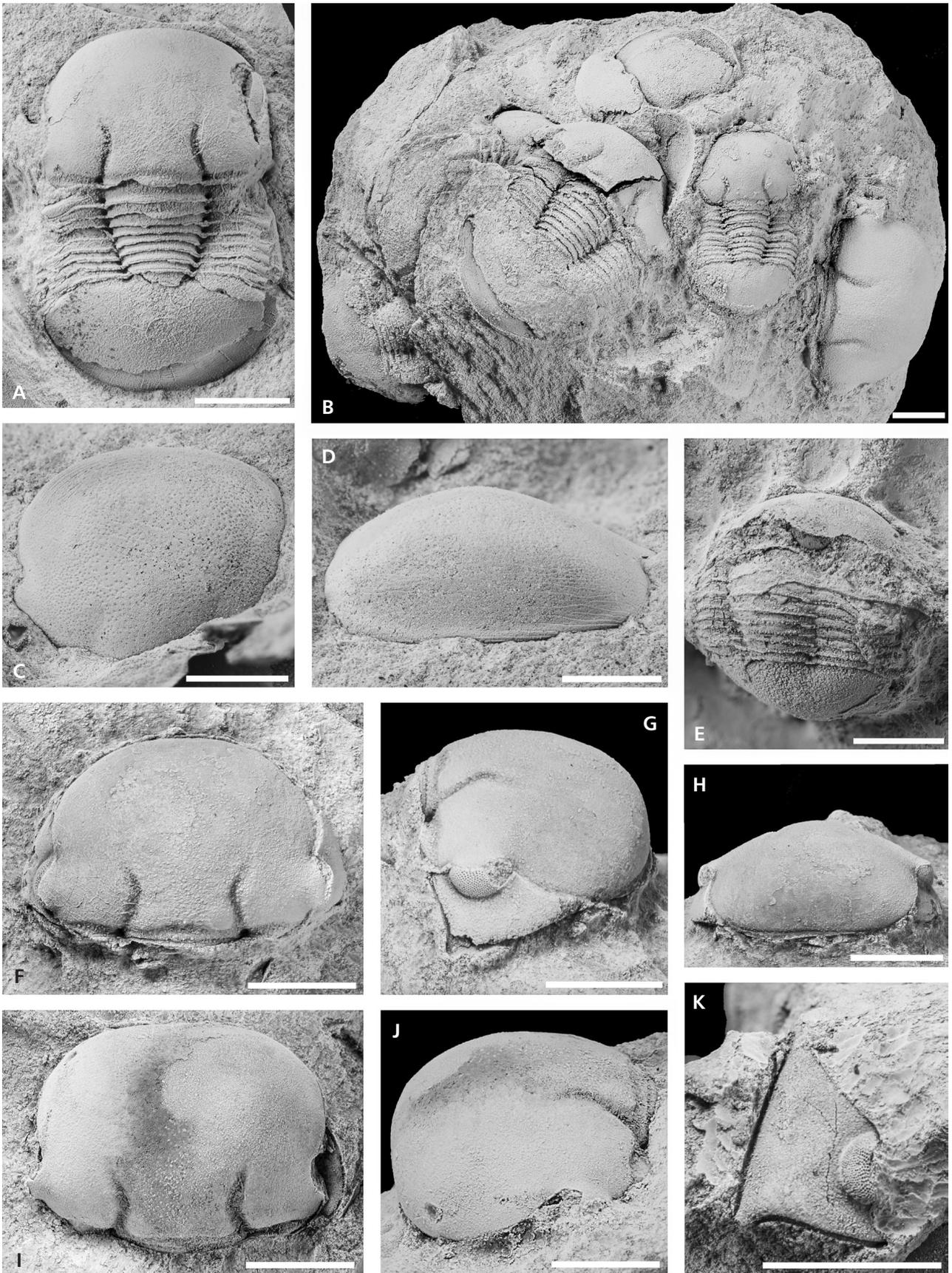
ize *O. hisingeri* and *‘I.’ oblitus* was incorrect. For the present work, the type material and the classical collections of both *O. hisingeri* and *‘I.’ oblitus* were revised and we consider this material to be conspecific. *‘Illiaenus’ oblitus* was erected and subsequently documented based on isolated sclerites from the upper part of the Králův Dvůr Formation, in dark grey shales and in the Perník (= Gingerbread) Bed (also named the ‘pelocarbonate’ bed), whereas *O. hisingeri* was defined and described based on prone or almost complete exoskeletons from underlying strata. However, the lectotype of *‘I.’ oblitus*, designated by Šnajdr (1957) and figured by Barrande (1872, pl. 15, figs 44, 45), Příbyl & Vaněk (1980, pl. 2, fig. 5), Shaw (2000, pl. 1, fig. 12) and Bruthansová (2003, fig. 12e), as well as other specimens figured as *‘I.’ oblitus* from the Perník Bed (e.g. Shaw 2000, pl. 1, figs 11, 15; Bruthansová, 2003, fig. 12c), are indistinguishable morphologically from little-deformed specimens of *O. hisingeri* specimens (compare with Šnajdr 1957, pl. 7, fig. 1; Šnajdr 1990, p. 141), including in the presence of eight thoracic segments. Characters of *O. hisingeri* previously listed as differences from *‘I.’ oblitus* (e.g. Barrande 1972, Šnajdr 1957, Hammann, 1992, Bruthansová 2003) are artefacts related to deformation. The palpebral lobes of isolated cranidia preserved as internal moulds seem smaller than when observed in complete cephalata (compare cephalata and isolated cranidia of *Vysocania* spp. or *Octillaenus maroccanus* in this work), the longer cranidia and pygidia are typical of sagittally stretched and flattened specimens (compare Bruthansová 2003, fig. 8a and Šnajdr 1957, pl. 7, fig. 1 or Shaw 2000, pl. 1, fig. 7), and the club-shaped glabella, anteriorly defined in some specimens (see Barrande 1852, pl. 29, figs 25, 27; Bruthansová 2003, fig. 8b) is due to the compaction of the cephalon against the rostral plate and the cephalic doublure. Undeformed and permineralized specimens assigned to *O. hisingeri* (e.g. Bruthansová 2003, fig. 8d, here figured in Fig. 9E) show the cephalic and pygidial outlines and pitting typical of specimens of *‘I.’ oblitus* from the Perník Bed. We therefore regard *‘I.’ oblitus* as a junior synonym of *O. hisingeri*.

***Octillaenus maroccanus* sp. nov.**

Figures 10, 11A–E

- 1985 *Stenopareia* sp. – Destombes *et al.*, p. 202.
 ? 2006 *Octillaenus* sp. – Destombes, pl. 40, fig. 7.

Figure 10. *Octillaenus maroccanus* sp. nov. from Upper Ktaoua Formation, Kralodvorian (ca. upper Katian), Zagora, Morocco; A – MGM 6898X (holotype), dorsal view (internal mould of an exoskeleton); B – MGUTAD 15702, internal moulds of several specimens; C, D – MGM 6901X, dorsal (C) and anterior (D) views (latex cast of an external mould of a cranidium); E – MGM 6897X, oblique view (internal mould of a partially enrolled exoskeleton); F, G, H – MGUTAD 15602, dorsal (F), lateral (G) and anterior (H) views (internal mould of a cephalon); I, J – MGM 6903X, dorsal (I) and lateral (J) views (internal moulds of a cranidium); K – MGM 6904X, dorsal view (internal mould of a librigena). All scale bars = 5 mm.



2010 *Stenopareia* aff. *oblita*. – Loi *et al.*, p. 350.

2011 *Stenopareia* aff. *oblita*. – Legrand, p. 291.

2014 *Octillaenus* sp. – Rábano *et al.*, p. 366.

Holotype. – Counterpart moulds of an exoskeleton lacking librigenae (MGM 6898Xa, 6898Xb; Fig. 10A).

Paratypes. – Four exoskeletons (MGM 6897X-1; MGUTAD 15701, 15702-1, 15702-2); two enrolled exoskeletons (MGM 6902X-2, 6912X); one cephalon (MGUTAD 15602); twenty-six cranidia (MGM 6895X-6907X, 6909X, 6914X-1, 6915X; MGUTAD 15602, 15700, 15702-3); one librigena (MGM 6904X-1); one rostral-plate (MGUTAD 15710-2); three thoracopyga (MGM 6895X-5, 6895X-6; MGUTAD 15702-4); six pygidia (MGM 6896X-2, 6896X-3, 6900X-1, 6910X, 6911X, 6914X-2; MGUTAD 15710).

Type horizon and locality. – Upper part of the Upper Ktaoua Formation; Kralodvorian (*ca.* Ka4, upper Katian), Bou Ingarf (30° 34' 22" N; 5° 34' 20" W), Zagora, Morocco.

Etymology. – After Morocco, from where the species is recorded.

Diagnosis. – Cranidium relatively short (sagittal length of 80% of width across palpebral lobes), extremely convex anteriorly in longitudinal profile; anterior margin slightly rounded; glabella about 45% of cranial width posteriorly; axial furrows moderately curved, attaining about 35% of cranial length; librigena subtrapezoidal with short genal spine; pygidium subelliptical to subpentagonal, with a low length/width ratio (about 58%); axis almost parallel sided, about 50% of pygidial length and 25% of maximum pygidial width.

Description. – Cephalon semicircular; cranidium subquad-rangular, strongly vaulted (sag. and tr.), overhanging anterior margin. Glabella corset-shaped, slightly higher than genal area in side view, width at posterior margin 45% of cranial width. Axial furrows slightly longer behind lunette than in front, dying out abruptly; lunette situated opposite anterior edge of palpebral lobe (γ). Occipital furrow broad (sag.), well defined on internal moulds, slightly convex anteriorly. Fixigenae broad (tr.), slightly narrower than glabella opposite lunette; posterior border furrow anteriorly arched, with depth and length similar to occipital furrow.

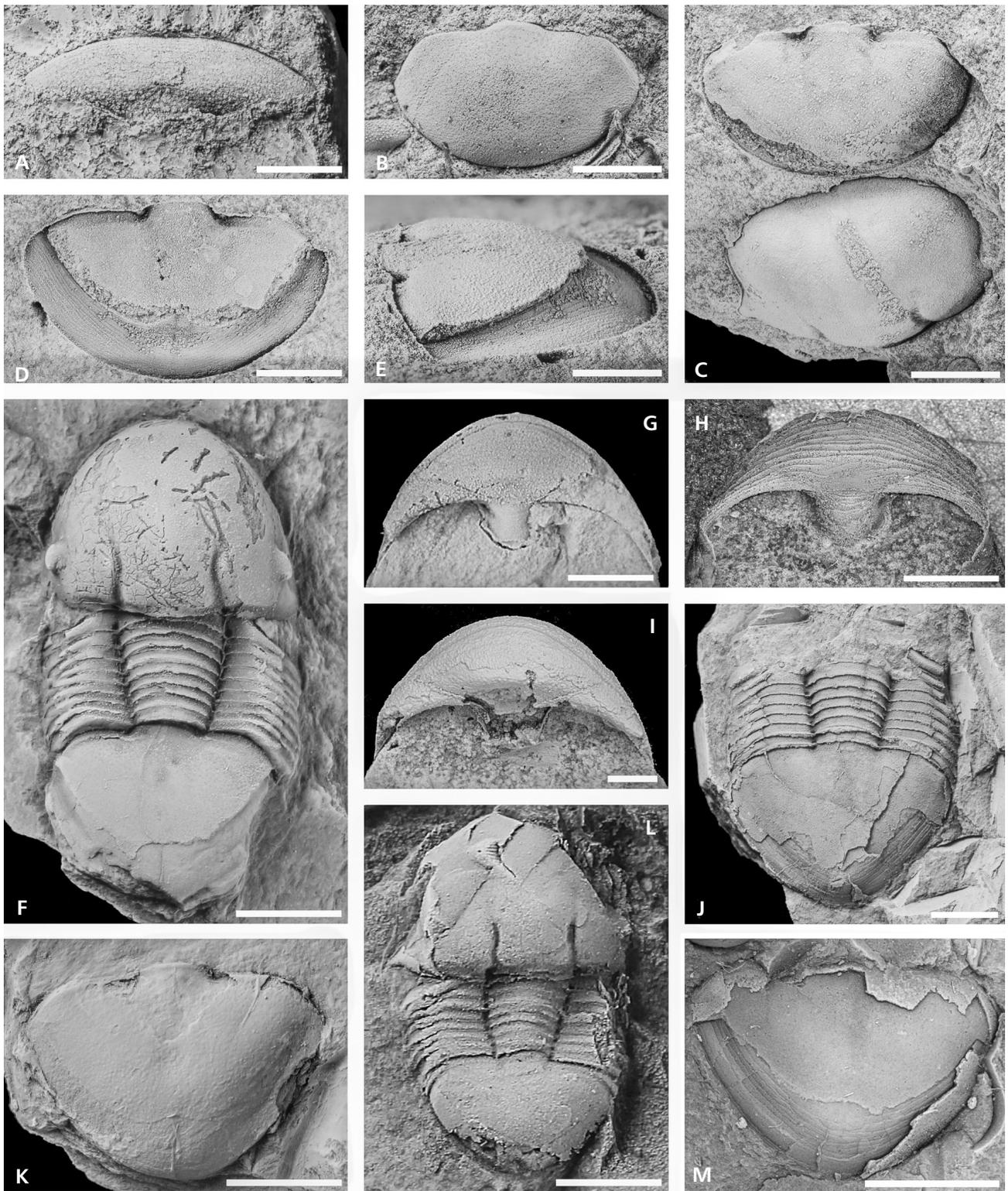
Posterior branch of facial suture very short (exsag.), exsagittal to convergent posteriorly; anterior branch divergent in posterior half, thereafter curving inwards (anterolateral cranial corners broadly rounded). Palpebral lobes semi-circular, corresponding to 17% of cephalic length, located posteriorly (cranial portion anterior to palpebral lobes corresponding to 66% of sagittal length); visual surface with dorsal lenses slightly larger than ventral ones. Librigenae subtrapezoidal, small, widest posteriorly, bearing a short, posterolaterally-directed genal spine; its external margin is continuous in outline with the librigenal lateral margin and forms an angle about 135° with the posterior librigenal border; librigenal doublure narrow (tr.), bearing about eight straight terrace ridges subparallel to the librigenal margin. Dorsal surface of cranidium with sculpture of homogeneous and dense pits, sparser on glabella than on fixigena; pits interspersed anteriorly with sinuous and discontinuous terrace ridges (about 15) running subparallel to anterior margin but seldom preserved; librigenal doublure with about eight continuous terrace ridges subparallel to outer margins.

Thoracic morphology like that of *Vysocania iberica*; exoskeletal pitting significantly finer than cranidium and pygidium.

Pygidium with pleural facet oriented at about 35–40° to sagittal line. Axial furrow not impressed; axis slightly convex (tr.), defined by forward deflection of anterior margin, change in convexity and muscle scars (rarely preserved). Dorsal surface except pleural facets with sculpture of pits as on cranidium, sparser on axis than on axial region, interspersed with sinuous and discontinuous terrace ridges abaxially and posteriorly; terrace ridges on pleural facets; doublure with about 15 uniformly spaced terrace ridges running subparallel to outer margin.

Remarks. – *Octillaenus maroccanus* differs from the type-species *O. hisingeri* in having shorter (exsag.), more strongly curved cephalic axial furrows, a more convex (sag.) cephalon anteriorly, and a slightly shorter (sag.) pygidium (these two last differences may be exaggerated by deformation, but are supported by undeformed specimens of *O. hisingeri*, *e.g.* Fig. 9E). Although we were unable to verify that the first thoracic segment in *O. maroccanus* is enlarged, as in *O. hisingeri*, one poorly preserved specimen suggests that it is (Fig. 10E). Although several specimens of *O. maroccanus* were prepared to show the hypostome, it remains unknown.

Figure 11. A–E – *Octillaenus maroccanus* sp. nov. from the Upper Ktaoua Formation, Kralodvorian (*ca.* upper Katian), Zagora, Morocco; A – MGUTAD 15710-2, ventral view (internal mould of a rostral plate); B – MGM 6900X-1b, dorsal view (latex cast of an external mould of a pygidium); C – MGM 6896X, dorsal and oblique views (internal moulds of a pygidium and a cranidium); D, E – MGUTAD 15710, dorsal (D) and lat-



eral (E) views (composite mould of a pygidium). • F–M – *Zetillaenus wahlenbergianus* (Barrande, 1852) from the Králův Dvůr Formation, Kralodvorian (upper Katian), Czech Republic; F – NMP L32898, dorsal view (internal mould of an exoskeleton); G – CGS-MS-1, ventral view (internal mould of a cephalic doublure and hypostome); H, I – NMP L32891, ventral view (latex cast of external mould, H, and incomplete internal mould, I, of a cephalon with hypostome); J – NMP L46558, dorsal view (internal mould of a thoracopygon); K – NMP 36675, dorsal view (internal mould of a pygidium); L – NMP L20412, dorsal view (internal mould of a meraspid exoskeleton, degree 8); M – NMP 40-99, dorsal view (composite mould of a pygidium). Scale bars = 5 mm except in I, M (= 2 mm).

Biostratigraphy, palaeobiogeography and evolution

Vysocania is an endemic genus from the high latitude peri-Gondwana realm (for Trilobita, the Dalmanitoidean realm, *sensu* Adrain *et al.* 2004), being recorded from the Czech Republic, Portugal, Spain and possibly France and restricted to the Berounian regional stage. The earliest known occurrence of the genus is that of *Vysocania moravecii* from the Sandbian (regional lower Berounian), in the Libeň Formation of the Czech Republic. This is the only documented occurrence in the lower Berounian, but it is likely that *Vysocania* is also represented among coeval Moroccan illaenid material from the Izzeguiene and Lower Ktaoua formations, which has been mentioned (*e.g.* Destombes *et al.* 1985, Horný 1997, Gutiérrez-Marco *et al.* 2003, Rábano *et al.* 2014) but not formally identified, described or figured. The youngest occurrence of *Vysocania* in the Czech Republic is that of *V. vaneki* from the upper Berounian Bohdalec Formation. In Ibero-Armorica, lower Berounian trilobite assemblages are characterized by low diversity (Henry 1980) and illaenids have not been documented in them. In Portugal and Spain, *Vysocania* first appears in the middle Berounian, in coeval units bounded by an oolitic ironstone horizon (Favaçal and Chôsavelha beds in Portugal; Young 1988 and Sá *et al.* 2005, respectively), represented by *V. iberica*, which ranges into the upper Berounian.

Vysocania is particularly dominant in basal ironstone horizons with low trilobite diversity. In the Czech Republic, the genus is present at the base of the Vinice Formation, in the Zdice-Nužice ‘Horizon’ where trilobites are rare, and in the low-diversity Karlík ‘Horizon’, at the base of the upper Berounian Bohdalec Formation, the type-horizon for *Vysocania vaneki*. In Portugal, *V. iberica* is dominant in the ironstone beds that mark the base of the middle Berounian Cabeço do Peão and Louredo formations (the Favaçal Bed) and in the upper Berounian–Kralodvorian Porto de Santa Anna Formation, where trilobites are extremely rare. Thus, species of *Vysocania* seem to have had a great capacity to adapt quickly to environmental change, with the competitive ability to occupy new biotopes. The abundance and broad geographic distribution of these species in the Berounian units of the high latitude peri-Gondwana realm lend them value as index fossils.

The study of abundant and well-preserved material of *Octillaenus marocanus*, together with the revision of the genus *Vysocania*, suggests that *Octillaenus* belongs to the same evolutionary lineage, developing through paedomorphic (neotenic) processes. *Vysocania* and *Octillaenus* share many of the same diagnostic features, and the only differences may be attributable to paedomorphism: eight instead of nine thoracic segments and a shorter pygidial doublure with a narrower and shorter

anteromedial deflection. Paedomorphism, in particular neoteny but also hypomorphosis (progenesis), is common among trilobites (*e.g.* Hupé 1953; McNamara 1981, 1986 and references therein; Crônier 2013 and references therein). Meraspid transitory pygidia of *Vysocania vaneki* from the Bohdalec Formation show that the doublure is shorter (sag.) in juveniles than in holaspides (compare Figs 3K–O). The same ontogenetic trend of an increase in the length of the pygidial doublure is shown by holaspides of *V. iberica*, with smaller specimens having shorter (sag.) doublures than larger ones (*e.g.* Hammann, 1976, pl. 3, fig 27a, and Fig. 4 herein). This trend may explain the typically shorter (sag.) pygidial doublure of *V. vaneki*, a species represented by small specimens, compared to the stratigraphically older species *V. panderi*. Alternatively, the shorter pygidial doublure of *V. vaneki* might be explained by an evolutionary trend in this character, similar differences in the length of the pygidial doublure being documented in other illaenid genera (*e.g.* *Stenopareia glaber* and the younger *S. linnarssoni*; compare Bruton & Owen 1988, figs 9d and 10a, b). The generally smaller size of *Octillaenus* species compared to those of *Vysocania*, the shorter pygidial doublure and its similarities with the youngest *Vysocania* species in the Czech Republic (*V. vaneki*; compare Figs 3M and 11D), the presence of one less thoracic segment and the stronger pitted sculpture are characters typical of *Vysocania* meraspides. In addition, in some *Vysocania* specimens (*e.g.* *V. iberica*, Figs 3F, 4I, 5C) the first thoracic segment seems slightly longer (exsag.) than the remaining segments (not significantly, tr.). Thus, the macropleural condition of *Octillaenus* may be another juvenile character retained from its presumed ancestor *Vysocania*. Apart from the number of the thoracic segments, it is extremely difficult to distinguish *Octillaenus* and *Vysocania*; cephalae and pygidia share the same dorsal features, and the rostral plate is similar.

Since relatively complete ontogenetic sequences are known for only a few members of the Illaenidae (*e.g.* Chatterton & Ludvigsen 1976, Chatterton 1980), it is difficult to use developmental data in resolving problems in the systematic relationships and phylogeny of the group. Furthermore, certain characters currently regarded as of taxonomic significance (*e.g.* the configuration of the pygidial doublure) may be more dependent on ontogenetic variability, being affected by paedomorphic processes, rather than being unique synapomorphies.

The morphological similarity between *Vysocania* and other *Stenopareia*-like trilobites is here interpreted as due to homoeomorphy. Another example of this phenomenon within the Illaenidae is *Eastonillaenus* Edgecombe & Webby, 2007 from the Katian (Eastonian of the Australian regional chart) of New South Wales. *Eastonillaenus* is very similar to *Vysocania* in the form of the cranidium, rostral plate and hypostome (see Edgecombe & Webby 2007,

figs 1c, 3b, d, f, g), but is clearly distinct in the thoracic and pygidial structure, having a very broad axis that gently tapers posteriorly. The similarity of both genera in ventral structure may also suggest that *Vysocania* is related to a geographically widely-dispersed clade of Illaenidae, including such genera as *Parillaenus*, *Eastonillaenus* and possibly *Zetillaenus*. Despite the previously documented differences between *Zetillaenus* (upper Katian) and *Vysocania* (Sandbian-lower Katian), both taxa have several characters in common, including not only the general dorsal morphology (e.g. axial furrows, occipital and fixigenal posterior furrows, genal spines, position of the eyes, facial sutures, axial width) but also some ventral features, namely the pygidial doublure and the hypostome (the only difference in the latter being the distinct shoulders in *Zetillaenus*).

After the works of Jaanusson (1954), Lane & Thomas (1983) and Whittington (1997), the need for a comprehensive revision of the Illaenidae incorporating both dorsal and ventral characters became evident. The lack of information on the ventral morphology of many illaenids is one of the major impediments to improve the revision of this family.

Conclusions

Vysocania is an independent genus distinguishable from *Stenopareia* and *Zetillaenus* and endemic to the high latitude peri-Gondwanan realm. Three previously named species are here included in *Vysocania*: ‘*Zbirovia*’ *vaneki* Šnajdr, 1958, the type-species, ‘*Dysplanus (Zetillaenus)*’ *ibericus* Hammann, 1976 and ‘*Illaenus*’ *panderi* Barrande, 1852. One new species is erected, *V. moravecii* from the Upper Ordovician of the Czech Republic. ‘*Illaenus*’ *oblitus* Barrande, 1872 is synonymized with *Octillaenus hisingeri* (Barrande, 1846), the type-species of *Octillaenus*, and the new species *O. maroccanus*, from the Upper Ordovician of Morocco, is erected. *Octillaenus* is considered to have arisen from *Vysocania* through pedomorphic (neotenic) processes. From an ecological point of view, species of *Vysocania* seem to have had good adaptability to environmental changes and the competitive ability to occupy new biotopes, as a result being particularly dominant in the low trilobite diversity palaeoenvironments recorded in Ordovician ironstone horizons.

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